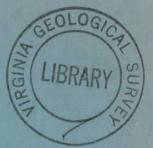
BUREAU OF ECONOMIC GEOLOGY The University of Texas Austin 12, Texas

JOHN T. LONSDALE, Director



Report of Investigations—No. 11

Correlation Between Surface and Subsurface Sections of the Ellenburger Group of Texas

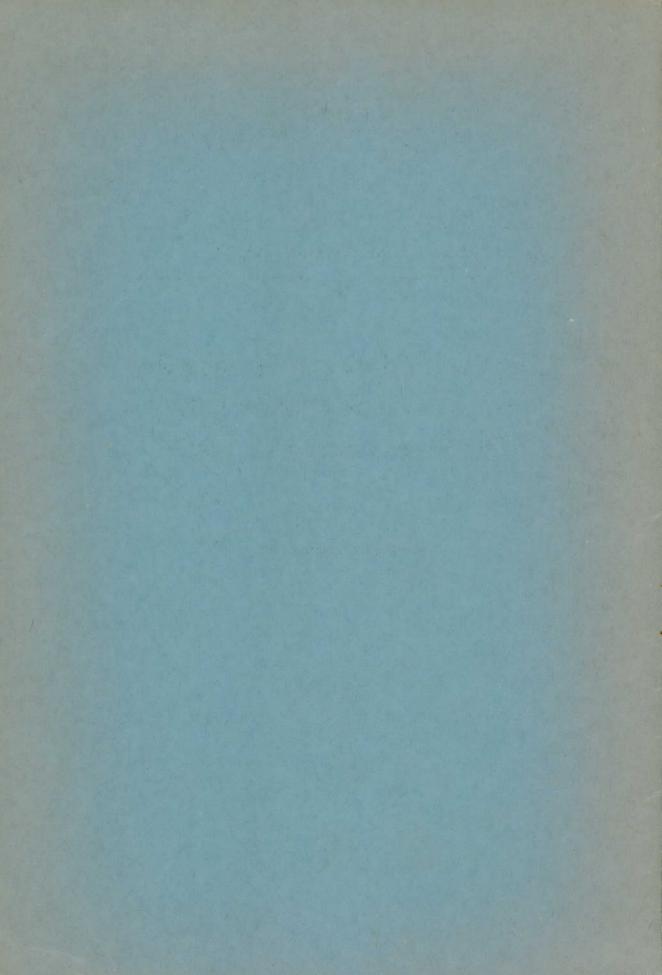
By

LEO HENDRICKS



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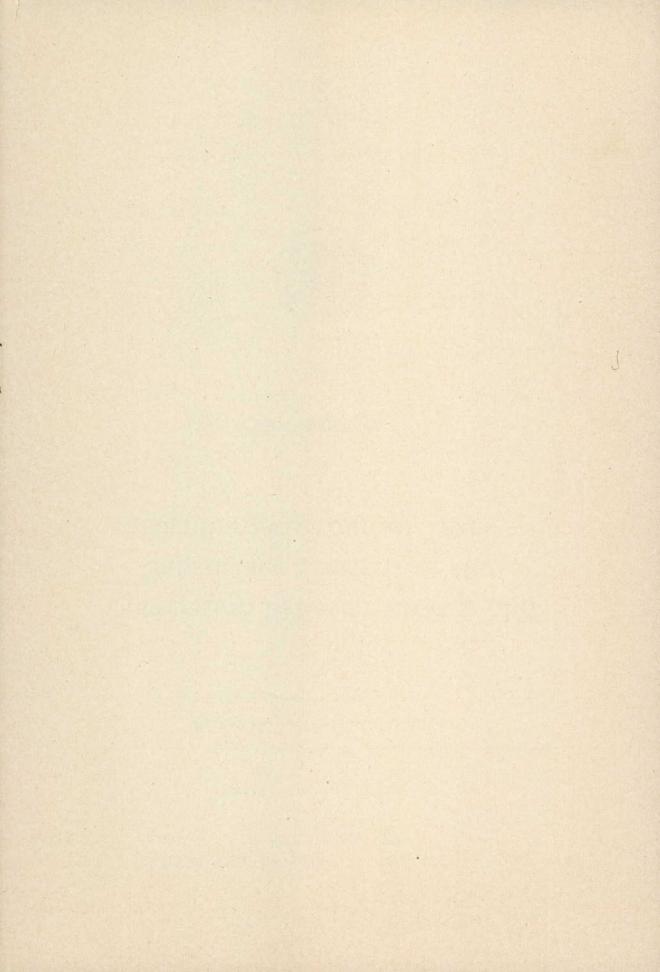
Correlation Between Surface and Subsurface Sections of the Ellenburger Group of Texas

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CORRELATION BETWEEN SURFACE AND SUBSURFACE SECTIONS OF THE ELLENBURGER GROUP OF TEXAS

Leo Hendricks

INTRODUCTION

The Ellenburger of Texas was first defined as a marine limestone formation of Cambrian and Ordovician age (Paige, 1912), but recently it has been subdivided

into several formations and the term Ellenburger given group status (Cloud and Barnes, 1948). The group forms an important unit in the geology of Texas, its

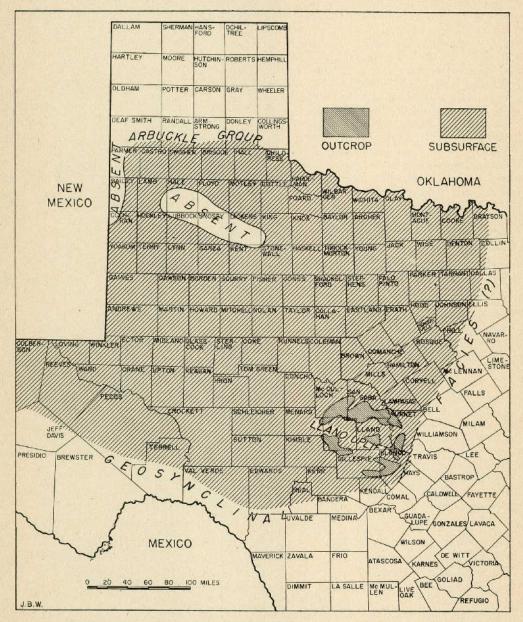


Fig. 1. County outline map showing approximate extent of Ellenburger in Texas.

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known extent in both surface and subsurface covering approximately one-half of the State (fig. 1). Consequently the Ellenburger has received the attention of many geologists over a period of more than forty years. The greatest amount of information concerning the Ellenburger has come from wells drilled in exploration for oil and gas. This information from the subsurface is the basis for present concepts of the lateral extent and regional changes in thickness and lithologic character of the group (Sellards, 1933b). The top of the group serves as an important key horizon for mapping structure in the subsurface of large parts of north, central, and southwest Texas (Sellards and Hendricks, 1946). Data from well samples and cores have been used to make qualitative subdivisions and correlations within the formation (Cole, 1942, p. 1398; Crowley and Hendricks, 1945, p. 413). No paleontological correlations are possible from well data because well samples and cores from the Ellenburger are practically barren of fossils.

Use of the term "Ellenburger" has some geographical as well as geological limitations. Rocks in the subsurface of northern Texas referred to as Ellenburger are identical with rocks in the subsurface of southern Oklahoma referred to the Arbuckle group (Decker, 1939). The two groups exhibit recognizable differences in their respective type areas, but the gradation from one to the other is so gradual that no welldefined transition zone can be recognized. Both the Arbuckle and the Ellenburger are assumed to lose their identity eastward in a geosynclinal clastic facies of Ordovician rocks (Sellards, 1938). Southward the Ellenburger possibly grades in part into the limestone formation of Lower Ordovician age recognized in the Marathon uplift as the Marathon formation (King, 1937). Westward a portion of the Ellenburger probably grades into the Lower Ordovician El Paso limestone (Richardson, 1909). Northwestward the group in part may grade into the Manitou formation of Colorado (Brainerd, Baldwin, and Keyte, 1933). Thus the rocks referred to as Ellenburger are largely a facies unit deposited in a great epicontinental sea.

The best evidence for stratigraphic correlation of the group has come from field study and interpretation of faunal evidence from the outcrop. Dake and Bridge

	The state of the s					
	Bear Spring area	Cherokee	Gorman and Tanyard	Warren Springs- Moore Hollow	Backbone Mountain area	Johnson City srea
1	070	1190	1457	1519	1467 - *	1096
Ellenburger group (total)	616	0711	1401	7101	. ±104T	10701
Honeveut formation (total)		142	325	453	402	629
Corman formation (total)	450	426	474	456	483	490
Calcitic facies	237	275	383	242	263	246
Dolomitic facies	213	151	16	214	220	244
Tanvard formation (total)	523	560	658	603	572*	657
Staendebach member (total)	229	300	456	353	481*	415
Calcitic facies	and a second second	176	137	121		33
Dolomitic facies	229	124	319	232	481*	382
Threadzill member (total)	294	260	202	250	16	242
Dolomitic facies		260	64	174	16	59
Calcitic facies	294		138	76		183

(1932) recognized several faunal zones of Cambrian and Ordovician age in the beds mapped as Ellenburger. In a recent cooperative project V. E. Barnes, of the Bureau of Economic Geology of The University of Texas, and P. E. Cloud, Jr., of the U. S. Geological Survey, (Cloud and Barnes, 1948) mapped portions of the Ellenburger outcrop in detail in the Llano uplift of Texas (Pl. I). A summary of their major concepts concerning the group follows.

The Ellenburger group is now limited to beds of Lower Ordovician age.

The Ellenburger is known to be present at the surface only in the Llano uplift region of central Texas, also known as the Central Mineral region. The group was studied and mapped in detail, and complete sections were measured in eight areas within the Llano uplift region (Pl. I). Reconnaissance studies were made throughout much of the outcrop area of the group, which surrounds and lies partly within the large topographic basin at the center of the Llano uplift. Three formations can be recognized as making up the group in the area of its outcrop.

The Tanyard, the oldest formation, is an inconsistent series of limestone and dolomite beds that averages 590 feet thick in seven measured sections. Changes in thickness are due chieffy to thinning from east to west. The limestones and dolomites characteristically have very abrupt lateral gradation from one to the other. Two members are recognized within the formation.

The Threadgill, lower member of the Tanyard, is predominantly dolomite but may be partly or all limestone, depending on facies. The member is essentially nonglauconitic, nonarevaceous, and without silt in the eastern side of the region, but silt is present in it on the western side. The member averages 234 fect thick in eight sections measured. Some bases for recognition in the field are diagnostic fossils, coarseness of grain of dolomite beds, typically nonporcelaneous chert.

The Staendebach, upper member of the Tanyard, generally is made up of a lower dolomite unit and an upper limestone unit but in places is all dolomite. The member averages 359 feet thick in seven sections measured. Field recognition is based largely on intermediate and fine grain size of the dolomite, the typically porcelaneous chert, and diagnostic fossils.

The Gorman formation lies on the Tanyard with some slight evidence of local disconformity. The formation consists generally of a lower unit of microgranular dolomite and an upper unit of limestone with dolomite occurring in the middle of this upper unit. The formation averages 463 feet thick in six sections measured where crosion had removed little or none of the rocks before later deposition. Recognition of the formation is based on diagnostic fossils, the usually very fine grain of the dolomites, the typically chalcedonic to porcelaneous chert, and the presence of sand grains occurring in zones.

The Honeycut formation, lying conformably on the Gorman, consists in its fullest known development of a lower unit of alternating limestone and dolomite, a middle unit of dolomite, and an upper unit of limestone. The dolomites are generally microgranular, and the cherts are chalcedonic to porcelaneous. Recognition depends on diagnostic fossils, generally lighter color of the dolomites, absence of sand except in the lower 50 feet. The formation was entirely removed by erosion before later deposition in the western half of the outcrop area, and it was differentially eroded over the remainder of the area. Thus no definitive upper limit can be placed on the formation from surface study. Its greatest known thickness at the surface is 679 feet.

Field evidence and succession of faunas indicate that deposition was almost if not entirely, continuous from time of deposition of the oldest to the youngest rocks in the Ellenburger group. Conglomerate at one locality and a slight irregularity of contact at another indicate a possible break in sedimentation between the Tanyard and the Gorman. Sediments apparently were rather pure, chemically precipitated, lime muds that became limestone or dolomite under dia-genetic processes. The Ellenburger in the area of its outcrop is remarkably free of terrigenous material, sand in the Gorman formation being, the most notable exception. The bulk of the evidence indicates that deposition occurred in a large, shallow-water area separated from land to the south and east by a deep trough, or by persistent currents that successfully withheld water-borne terrigenous material. Changes in faunal make-up from one level to another possibly were due to regional changes in ocean depth, salinity, temperature, or other conditions affecting life in the sea.

The Ellenburger lies with apparent conformity on the Upper Cambrian in the western part of the Llano uplift region, but irregularity of contact indicates it may be unconformable with the Upper Cambrian in the eastern part of the region. Throughout the Llano region the Ellenburger has a compound unconformity at its top, having been reached by erosion in some or all of its extent by at least eight cycles of erosion, each followed by overlapping deposition. Differential erosion has brought about a thinning of the Ellenburger westward across the area of its outcrop (Table 1, p. 6). Devonian and younger formations rest to west. The Honeycnt formation has all been removed west of the longitude of western San Saba County.

The Ellenburger contains beds of earliest Ordovician age, being apparently conformable, in at least a part of its occurrence, with Upper Cambrian. The group does not contain beds equivalent to youngest Lower Ordovician in its outcrop area. Lower Ordovician beds younger than Ellenburger are present at the surface in the Marathon, Van Horn, and El Paso regions of west Texas, in the Wichita and Arbuckle Mountains of Oklahoma, and in the Ozark region of Missouri and Arkansas. Regional correlation with these areas is shown on Plate VI.

The above summary presents briefly the nature and stratigraphic significance of the Ellenburger as now defined in the area of its outcrop. In its subsurface occurrence the Ellenburger has been defined as including beds up to the base of the Simpson formation of Middle Ordovician age (Sellards, 1933a). Thus the Ellenburger in a portion of its occurrence may contain some beds younger than any occurring in the Honeycut formation in its thickest development. The youngest known beds assigned to the Honeycut formation are overlain by rocks of Devonian age. It seems most likely that the hiatus of the unconformity includes beds present underneath Middle Ordovician in west and north Texas. Whether those Lower Ordovician beds younger than Honeycut as defined should be included in the Honeycut becomes a matter of judgment based on evidence from many sections. Especially, subsur-face sections of north Texas should be compared with sampled sections of Lower Ordovician rocks present in the Arbuckle and Wichita Mountains of Oklahoma.

An understanding of the occurrence of the Ellenburger formations in the subsurface involves the use of dependable correlative criteria common to both outcrop and subsurface data. Since faunal evidence obtained so far in well data is too rare to be of use, the physical or qualitative characteristics of the Ellenburger subdivisions offer the only recourse for correlative data to be found in both outcrop and well samples. The major lithologic characteristics may serve to distinguish the group as a whole from other rocks but do not offer evidence for widespread correlation within the group. Cloud and Barnes found that grain size in the dolomites was a general guide for recognizing some subdivisions, and this is true to some extent in more widespread subsurface correlation. The alternation of limestone and dolomite has no stratigraphic significance, since one grades laterally into the other in an unpredictable manner.

The insoluble residues of the Ellenburger offer the best qualitative evidence found to date for recognizing subdivisions within the group. The possibility of subdividing the Ellenburger in certain areas of its subsurface extent on the basis of insoluble residues alone has been demonstrated (Cole, 1942; Crowley and Hendricks, 1945). Cloud and Barnes noted characteristics of the chert impurities in the limestones and dolomites on the outcrop peculiar to certain formations. The characteristics are evident in both the embedded chert and the chert formed at the surface under the processes of erosion. It will be shown that samples taken from surface sections and digested in acid will yield residues that show a qualitative variance from one level of the Ellenburger to another. The change in residue quality does not coincide with formational boundaries in every case, but the points of change are sufficiently consistent that the residues offer a clue to formation identification, if not exact definition. The chief asset of the residues is the fact that variations in residue quality can occur with age despite little or no change in major type of sedimentation. The chief weakness of the insoluble residues as an aid to stratigraphic study of the Ellenburger is the fact that residue quality is an expression of facies and can change laterally with change in sedimentary conditions affecting deposition of insoluble material.

The term "insoluble residue" is here used in the same sense as others have used it (McQueen, 1931; Ireland, 1936). The samples were treated with dilute hydrochloric acid or, in some cases, with dilute acetic acid. The residual material is therefore known to be insoluble only in those acids.

One of the objectives of field work on the Ellenburger was to obtain sets of samples from measured sections whose insoluble residues could be compared with those from subsurface sections. Therefore Cloud and Barnes selected and marked sections that offered the best opportunity for obtaining samples from beds in place. The very careful and painstaking field work of Cloud, Barnes, and their assistants, principally L. E. Warren and R. L. Heller, has made it possible to collect dependable surface samples for comparison with subsurface samples. It is with great pleasure that this opportunity is used to express appreciation for the excellent manner in which they clarified the field relationships of the various sections. Grateful acknowledgment is made for their generous cooperation in gathering samples from the field for comparison with samples from the subsurface.

Other sources of help on the subsurface study of the Ellenburger include various oil companies and geologists operating in north and west Texas. Their generous help in making samples and residues available for study made the project possible.

The work of investigation has been carried on under the auspices of the Bureau of Economic Geology of The University of Texas, and the writer wishes to acknowledge the great value of the assistance and encouragement received from the Director of the Bureau.

ACQUISITION AND STUDY OF RESIDUES FROM SURFACE SECTIONS

The insoluble residue characteristics of each of the Ellenburger formations presented here are revealed in samples from measured sections in each of the areas mapped in detail by Cloud and Barnes (Pl. I). The samples are from six composite sections, representing the total measureable thickness for each of the six areas, and from one incomplete section. The sections are designated as follows:

- 1. Gorman Falls--Tanyard
- 2. Cherokee Creek
- 3. Warren Springs-Moore Hollow
- 4. Backbone Mountain
- 5. Johnson City
- 6. Threadgill Ćreek

7. Llano River

Sampling the sections was done in part by Cloud, Barnes, and assistants and in part by the writer, assisted by L. E. Warren. All the sampling was based on measurements by Cloud and Barnes. Preparation of all residues was under the supervision of the writer.

In sampling the Ellenburger, fresh chips were broken from beds at as short vertical intervals as exposures would allow. Chips. from a 5-foot interval were combined to form one sample. Different types of rock occurring in a 5-foot interval were represented in their proper proportion so far as sight judgment could allow. Each sample was crushed in a small jaw crusher and a thorough mixing of fragments from each chip obtained. A uniform portion of each crushed sample, averaging about 30 grams of the material, "was used in making residues. The fines were not screened from the crushed material, but the rock flour resulting from the crushing was removed by washing and decanting before adding the acid. Solution was carried out in 400-ml. pyrex beakers, using approximately 12 percent hydrochloric acid, or approximately 10 percent acetic acid. After complete digestion of the soluble rock the clay-size particles were removed from the residue by washing and decanting. The remaining residue was dried on an electric hot plate and bottled.

This method of collecting and preparing surface samples for study was followed because experience in using residues from the Ellenburger as an aid in stratigraphic study has shown that the mass characteristic of the residues is the most informative. Few, if any, residual materials are individually diagnostic of a formation. Hence in sampling for residue study, as much as possible of the vertical extent of the formation must be included. Trench sampling would be ideal but is impractical. Spot sampling at intervals of a few feet will not catch enough of the mass characteristics.

The residues were examined under a binocular microscope, using a magnification of 15.6 diameters for the most part. Light source was a fluorescent lamp with a daylight-type bulb. The data considered were almost entirely qualitative, being quantitative only to the extent of sight comparison of quantities of materials.

Residues obtained from the Ellenburger by treating with hydrochloric acid are made up of four broad types of materialchert, non-clastic quartz, clastics, and accessory constituents. Each of these types occurs in a variety of forms and combinations, thus producing the qualitative variations that are significant in stratigraphic interpretation. Chert is the most significant material, occurring in a wide variety of textures, structures, and degrees of transparency. The distinction between chert and quartz is somewhat arbitrary, since all types of chert are varieties of the mineral quartz. In describing the residues the term quartz is reserved for material with the fracture, luster, and crystal habit of the mineral. By assigning symbols, either colored or drawn, to represent the various materials in residues, the succession found in a section can be presented on a log strip. Correlation between sections can best be accomplished by matching plotted logs.

Samples from limestone beds in the Ellenburger may be dissolved in acetic acid, in which the dolomite is insoluble. Samples from the limestone portions of several sections, both surface and well sections, have been treated with acetic acid for comparison with residues from hydrochloric acid. Few differences have been discovered between the two types of residues. The acetic acid residues from surface sections have yielded a few conodonts, but they are much too rare to be of use in stratigraphic work. The effervescence from acetic acid is less violent, thus leaving more of the delicate clay and fine clastic aggregates intact. This particular quality of the residues is an aid in distinguishing fine-grained, gray Ellenburger limestone from similar Mississippian limestone in the subsurface of north Texas. Soft shale fragments preserved in the acetic acid residues from the Mississippian limestone contain fossil imprints while those from the Ellenburger limestone do not. Although very closely related to the major sediments as to time of deposition, they may have an entirely independent source. On the other hand, the various types of chert found in the residues represent deposition from solution, just as the carbonates do. Part of the cherts appear to be syngenetic in relation to the enclosing rock, and part have a secondary or epigenetic appearance. The source of the syngenetic cherts may be the same as, or very closely related to, the source of the major sediments. The epigenetic cherts may be the result of alteration of syngenetically deposited siliceous material, or their silica may have been brought in by circulating ground waters after the deposition of the carbonates and their accompanying syngenetic materials.

Table 2. Genetic classification of residues.¹

ALLOGENIC	AUTHIGENIC	
	Syngenetic	Epigenetic
Silt	Chert (segregated)	Chert (interstitial)
Sand	Shale partings	Quartz .
Shale fragments	Clav masses	 Siliceous onlite and onliths
Mineral fragments	Pyrite (crystalfine)	Pyrite (interstitial)
(e.g., mica)	Clauconite	Anbydrite
Siliceous spicules	Sedimentary feldspar, mica	Replaced fossils
· · · · ·	Fossil fragments	

⁴Ireland. (1936) has given a similar classification of residues from lower Paleozoic formations of Oklaboma.

acetic acid residues do not seem to offer much additional aid in the area and portions of the section where they have been used, the method has not been given an exhaustive trial. In other areas and portions of the section positive results may be obtained.

EVALUATION OF RESIDUES AS AN AID IN ELLENBURGER STRATIGRAPHY

The insoluble residues obtained from the Ellenburger deposits represent the accessory constituents in the main body of rock. They are the products of minor sources of deposits which were greatly overshadowed by the carbonate deposits. The degree of relationship between the sources of these major and minor deposits may vary widely. On the one hand, sand grains and shale fragments in the residues represent mechanical deposition in the midst of a chiefly precipitated sediment and are syngenetic deposits in relation to the enclosing rock. Although they are

It is conceivable that minor sources of sediments, such as furnished the insoluble materials, may be modified or completely disrupted by a change in conditions too small to be reflected in the major portion of the sediments. For example, the failure of some transporting agent that had been depositing scattered sand grains in a calcareous deposit would not necessarily affect the deposition of limestone. Likewise, the modification of chert formation in minor amounts would not necessarily affect the carbonate deposits. However, any change in conditions that modified the type of insoluble accessory materials being deposited would be evident in the insoluble residues. Thus the residues may in some cases afford a more sensitive tool for detecting changes in conditions than the bulk of the sediments.

Accessory materials undoubtedly exhibit lateral gradation from one type to another during the same time interval, varying with the numerous local conditions influencing sedimentation in a basin. Some accessory materials, however, appear to be the result of regional influences and carry through minor changes in local conditions. Those changes in residues due to change in local conditions may be used for local correlation only. A more widespread change due to modification of regional influences may be used for wider correlation. Average temperature and solute content of the water, for examplé, are regional conditions subject to modification that may affect the accessory materials throughout a sedimentary basin.

In using insoluble residues as an aid in stratigraphic study of the Ellenburger each type of residue material has been carefully evaluated and is relied upon in correlation only within the probable limits of its dependability. The syngenetic cherts are the most widespread of the insolubles and are the most thoroughly disseminated through the group. This, perhaps, is due to their being rather closely related in method of deposition to the major sediments, which indicates that they are the result of regional influences and not local conditions. By the nature of their occurrence, then, the syngenetic cherts are the most important and diagnostic of the residue materials.

Other significant materials in the residues may be of diagnostic value but to a lesser degree than the syngenetic cherts. Any of the mechanically deposited materials such as sand grains, shale inclusions, and silty material may occur in zones that can be correlated locally—within the limits of an oil field, for example. Silicified ooliths and oolitic cherts are more likely to be due to specialized local conditions than to widespread influences.

RECOGNITION OF THE ELLENBURGER FROM STUDY OF SAMPLES

Gross lithology has long been a criterion for recognizing Ellenburger rocks from samples. Typical Ellenburger limestones have a high degree of purity, light color, sublithographic texture, and complete absence of any direct evidence of organic origin. Fracture of the limestone is sharp, with a minimum of powdering along the edges of small chips in crushed samples. Typical Ellenburger dolomites have a well developed sucrose texture, are generally light in color, and also exhibit a high degree of purity. These characteristics usually serve to distinguish the group from all rocks found lying immediately above the Ellenburger in the outcrop region. Dolomite in the Edwards formation of the Lower Cretaceous is the only similar rock in the area.

Recognition of the base of the Ellenburger from samples can be quite difficult where similar lithologies are in contact at the systemic boundary. In the eastern and southeastern portion of the outcrop area Ellenburger dolomites rest upon Wilberns. dolomites, as illustrated by the Johnson City section and the Gorman Falls-Tanyard section (Pl. II). The Wilberns dolomite is usually finer grained than the Ellenburger dolomite. Stratigraphic recognition is aided by a study of insoluble residues. Lowermost Ellenburger dolomites are characteristically lower in residue content, with the residues being marked by the presence of granulated chert and waxy shale flakes. Wilberns dolomites usually furnish more residue, with granular chert predominating and drusy quartz common.

In the western part of the outcrop area silty, somewhat impure Ellenburger limestones containing some glauconite rest on similar limestones in the Wilberns formation, as shown by the Llano River section (Pl, II). Here the texture of the limestones and the glauconite content furnish clues to stratigraphic identification. The Wilberns limestones are typically granular, in contrast to the extremely fine to sublithographic texture of the Ellenburger limestones, and contain a greater abundance of glauconite. Comparative abundance of glauconite is most easily observed in the residues from the limestones.

Residues alone are not a dependable source of evidence for identifying the Wilberns-Ellenburger contact throughout the outcrop area. Ellenburger-type residues, dominated by cherts, and Wilberns-type residues, high in silty, argillaceous material, can be recognized in samples from those formations throughout the outcrop area. But in the eastern sections Ellenburger-type residues extend down into the Wilberns dolomites, and in the western sections Wilberns-type residues extend up into the Ellenburger. In the eastern sections the boundary can be determined fairly closely from the residues because of a change in the nature of the chert, as noted above. In the western sections, however, no widely dependable diagnostic change occurs in the silty, argillaceous residues present both above and below the boundary. Increase in the amount of glauconite is an indication of Wilberns age. This, together with the granular texture of the limestones, seems to be the best criterion for distinguishing Ellenburger from Wilberns in samples from the western portion of the outcrop area.

The transgression of silty, argillaceous residue material across time lines as established by fossil evidence is shown on Plate II. The insoluble material is typical of Wilberns limestones but does occur in Wilberns dolomites; for example, in the Warren Springs-Moore Hollow section the material is present in the lower portion of the Pedernales dolomite member. Farther west, in the Threadgill Creek section, the lower half of the Threadgill member of the Tanyard formation yields silty, argillaceous residues. Still farther west, in the Llano River section, the material occurs throughout the Threadgill. The upper limit of the occurrence of silty, argillaceous residues in the Cambrian-Ordovician sections is presented as a transgressing but recognizable boundary on the correlation chart of Plate II.

RECOGNITION OF FORMATIONS IN ELLENBURGER GROUP

Recognition of the formations in the Ellenburger group from a study of systematically sampled sections must be based on rock characteristics alone since no microfaunal evidence is present. The rock characteristics considered in the Ellenburger samples consisted of the lithology and texture of the original rock and the various types of materials present in insoluble residues from the samples. Standard lithologic logs were prepared of all composite sections that were measured and sampled. In addition, highly detailed logs of the residues from each section were made. Comparison of the lithologic and residue logs from all the surface sections demonstrated that recognition of the Ellenburger formations from sample logs is possible to a fair degree of accuracy.

The standard lithologic logs showing texture of the dolomites can serve only as general guides when used alone. The best evidence for recognition of formations is found in the combinations of residue associations shown on the residue logs. The lithologic logs can be an aid in deciding exact points when used with the residue logs, for it often seems logical to place a formation top where a change in lithology or texture occurs.

Careful study of the highly detailed residue logs showed that the residue materials most pertinent to the recognition of the formations can be classified as follows:

> Non-granular chert Granulated chert Non-clastic quartz Sand Silt Argillaceous material Glauconite

Consequently a set of black and white and color symbols was devised for showing on logs the occurrence of these materials in the surface sections.

Non-granular chert is that in which no grain could be distinguished under low magnification. Chalky, earthy textured cherts, porcelaneous cherts, and intermediate types bordering on granular were usually logged as non-granular. Those logged as granular showed a definite grain on the broken surfaces. The granulated chert occurs in very rough irregular fragments that have the appearance of being aggregates of irregular chert grains.

The non-clastic quartz includes fragments of the mineral quartz that are dolomoldic, drusy, or show other evidence of having been deposited in place in the rock. X

The sand grains are rounded to subrounded, usually frosted, quartz fragments. The silt is made up of silt-size siliceous grains which appear to be clastic in origin. Some of the material, however, may be the result of precipitation and growth of small grains in place. The silt occurs both free and embedded in argillaceous material.

The argillaceous material includes flakes ' of shale, clay-like masses, and spongy, porous masses not properly described as either clay or shale.

For detailed discussion of residue material, see page 18.

TANYARD FORMATION

The outstanding sample characteristic of the Tanyard formation in the standard section area of Tanyard and Gorman Falls (Pl. 11) is the predominance of granular

chert in the residues of all but the uppermost beds of the formation (sample descriptions, p. 36). Non-granular chert ranging from smooth, conchoidal to chalky in appearance may be present but not as abundantly or persistently as granular chert. The granular chert from the dolomitic facies of the Tanyard is characteristically dolomoldic. Drusy quartz is fairly common in the formation. The chert and quartz of the formation are largely interstitial in occurrence below the weathered surface, although large masses of chert form at the surface as a result of concentration under weathering processes. Thus, crushed samples of Tanyard rocks do not commonly show free fragments of chert, even though the percentage of insoluble residue may be high. Oolitic chert, both granular and non-granular, appears sporadically in the Tanyard section.

In the eastern sections the Threadgill member of the Tanyard formation can be distinguished on the logs from the Staendebach. The Threadgill member contains the special type of granular chert best described as granulated chert. The chert appears as very rough, highly granulated fragments, added to the other chert materials common to the Tanyard. In some samples the granulated chert becomes the most abundant material. Green and red waxy shale flakes also appear more consistently in Threadgill residues than in Staendebach residues. Threadgill samples average a noticeably smaller percentage of residue material than either Staendebach or Wilberns dolomite samples.

A marked change occurs in the nature of the residues from Tanyard samples in moving westward across the outcrop area. The residue log of the Threadgill Creek section, Gillespie County (Pl. II), shows normal Tanyard-type residues from the upper half of the member, but the lower half has the argillaceous, silty residue material similar to residues from the underlying Wilberns. In the Llano River section, located still farther west in Mason County, the silty, argillaceous residues have replaced cherty residues through all of the Threadgill and have transgressed slightly into the base of the Staendebach. Most of the Staendebach section has normal Tanyard residues, with one notable feature: the granulated chert typical of Threadgill in

the eastern sections is present in the Staendebach in the Llano River section. Thus, the upper Tanyard can be recognized from residue samples in the western outcrops, but the lower portion is better distinguished by the subgranular texture of the limestones, as noted previously.

The uppermost beds of the Tanyard, averaging about 100 feet in the outcrop sections, do not furnish residues that are typical of the remainder of the formation. The cherts in the upper beds are largely non-granular and are similar in nature to those from the overlying Gorman formation. The chief distinction between the residues of the upper beds and those from the Gorman is the absence of sand in the Tanyard residues. This criterion appears dependable throughout the outcrop area.

It will be noted on the lithologic logs (Pl. II) that the texture of the Tanyard dolomites ranges from coarse to fine, with coarse and medium predominating. The coarsest dolomites are found in the dolomitic facies of the Threadgill member, with the Staendebach dolomite grain averaging somewhat smaller. The Tanyard limestones all have a very smooth fracture and are subgranular. In the eastern sections the limestones are pure and very light in color. In the western sections the lower limestones become grayish and furnish silty, argillaceous residues.

GORMAN FORMATION

Residue logs of the Gorman formation outcrop sections (Pl. II) show the chief constituents to be non-granular chert and (Sample descriptions of type secsand. tion, p. 26.) Subordinate materials are granular chert, quartz, shale, and other argiflaceous material. The non-granular chert consists largely of smooth opaque and smooth translucent varieties but includes some chalky and subporcelaneous. The chert for the most part is not interstitial, and free fragments are found in the crushed samples of the rock. The chert occurs in thin seams, beds, and irregular nodules in the limestones and dolomites and because of its manner of occurrence is not typically dolomoldic. Zones of oolitic chert are present. Near the middle of the formation some of the chert usually contains embedded sand grains. The sand in the Gorman

residues is composed of pure, frosted quartz grains, and the larger grains are rounded. The sand occurs both as scattered, individual grains and as sandy zones in the formations. Grain sizes in the sandy zones show poor sorting, ranging from fine to medium coarse.

The dolomites of the Gorman are persistently finer in texture than the Tanyard dolomites in the eastern Ellenburger sections. Grain sizes are fine to very fine. However, in the Llano River section the lower dolomites of the Gorman are coarse in texture.

The occurrence of sand together with the grain size of the dolomite in the eastern area furnish the criteria for determining both the base and the top of the Gorman from samples. The overlying Honeycut contains very little sand except in the lower 50 feet of beds.

HONEYCUT FORMATION

Residues from the Honeycut formation are composed largely of non-granular chert, similar to that in the Gorman. (Sample descriptions from Gorman Falls-Tanyard section, p. 23.) In general, there is a slightly smaller percentage of insoluble material in the Honeycut than in the Gorman. Scarcity of sand in the Honeycut residues, however, is the most noticeable difference in surface sample residues from the two formations. Thus, recognition of the Honevcut from residues must be based largely on the predominance of non-granular chert, with the absence of the sand that characterizes the Gorman. Dolomites of the Honeycut are similar to those of the Gorman but show a tendency to be finer textured. Most of the microgranular dolomite of the Ellenburger group is found in the Honeycut samples.

POINTS OF CORRELATION ON THE ELLEN-BURGER SURFACE SECTION SAMPLE LOGS

The type section for the Ellenburger group, as designated by Cloud and Barnes (1948), is the composite section made up of their Gorman Falls, Spicewood Creek, and Tanyard sections. They designated the composite section made up of their Warren Springs and Moore Hollow sections as a standard section for the Ellenburger. The logs of the two composite sections therefore are type logs to which other logs should be referred. Several correlative points are evident on the two logs that can be recognized on the logs of the other surface sections.

The most marked and persistent feature of the logs is the change on the residue logs from non-granular cherts above to granular cherts below at a level near the top of the Tanyard formation. The level marked by the change serves as a dependable correlative horizon for relating the sections on the basis of sample evidence. When the logs are leveled on that horizon, the tops of the Tanyard as determined by field and fossil evidence are not far from alignment. Thus it is felt that any section of the Ellenburger including this portion of the section can be closely correlated with the type Ellenburger section.

The presence of sand in the Gorman and in the base of the Honeycut and the absence of sand in the Tanyard are other characteristics appearing on all residue logs of those formations. Hence, the lowest occurrence of sand makes a fairly good correlative point, marking approximately the base of the Gorman.

A granulated chert zone is present in the Tanyard on all residue logs showing the complete formation. The granulated chert appears in the Threadgill member in the eastern sections and in the lower part of the Staendebach in the westernmost section.

The granulated chert zone is a good correlative marker in the facies of the Ellenburger represented by the castern sections but indicates somewhat younger beds with changing facies westward.

A sandy chert zone is typically present near the middle of the Gorman formation. The chert is usually free of sand even though sand is abundant in the residue. But in this particlar zone many of the sand grains are embedded in the chert.

The horizons present on the residue logs listed above furnish the means for correlating the logs of the sections and are a valuable aid in recognizing formations from sample information. No distinctive feature of the gross lithology occurring persistently at a given level is evident from the sample study.

PROBLEMS IN RECOONIZING ELLENBURGER FORMATIONS IN THE SUBSURFACE

The subdivision of the Ellenburger into formations rests primarily on faunal evidence. But in choosing formation boundarics Cloud and Barnes used the best physical breaks available in the rather homogeneous group of beds. Thus the formations are essentially time-rock units with fairly mappable boundaries marked by changes in such physical characters as grain size of the dolomite, occurrence of sand, and nature of contained chert. When physical evidence fails, fossils can be sought to help establish formation boundaries in surface mapping. Each formation carries a definite age meaning.

Recognition of formations from subsurface samples must be based on physical evidence alone. The point of greatest concern is whether or not the physical evidence available continues to have the same age significance that it has at the surface. Evidence from the surface section sample logs supports the idea that some of the physical evidence usable in correlating the sections does have age significance. The most dependable horizon recognizable in the samples is the level marked by the change from non-granular to granular chert in the residues. This change occurs consistently near the top of the Tanyard formation, regardless of the location of the section. Other marked zones; such as the granulated chert zone of the Threadgill member and the sandy chert zone of the Gorman formation, appear to have consistent age significance over a large area in the eastern portion of the outcrop area. In a westerly direction the granulated chert zone to some extent appears to migrate up the section across time lines. It is in the eastern portion of the outcrop area that the type and standard sections are located. Therefore, the succession as developed in the eastern sections. may be thought of as representing typical Ellenburger facies. Sample log evidence indicates that where typical Ellenburger facies is developed the physical evidence has age significance, and the time-rock unit type of formation can be recognized on the basis of that evidence.

CORRELATION OF ELLENBURGER SUB-SURFACE SECTIONS

CENTRAL TEXAS TO NORTH TEXAS

Hundreds of wells have penetrated a portion of the Ellenburger section in the subsurface of north-central and north Texas. A few wells have drilled through the group, and many have drilled several hundred feet. Enough sections are available for general correlation of the Ellenburger beds of the area with the formations defined at the surface.

Plate III presents the plotted logs and stratigraphic correlation chart for a series of sections on a traverse from the southern side of the Llano uplift northward to northern Archer County. All formations described at the surface can be recognized along the traverse at least as far north as Erath and Somervell counties. In northern Archer County the Tanyard-Gorman boundary can be determined, but available criteria fail to distinguish the Gorman-Honeycut boundary or to establish clearly the base of the Tanyard. Residues from the entire upper Ellenburger section become notably arenaceous in extreme north Texas. It is felt that the subsurface Ellenburger sections in this area can best be correlated with surface sections exposed in the Wichita Mountains of Oklahoma.

Plate IV presents the sections and their correlations on a traverse from Fisher County eastward to Somervell County.

The well sections of north Texas have been correlated with the surface sections very largely on the basis of evidence on the residue logs. The best control point for correlating the sections is again the contact of the granular and non-granular chert zones, as was true with the surface sections alone. This characteristic change in chert type has been observed in all the wells investigated in north Texas that contain that particular portion of the section. Because of its persistent appearance near the same stratigraphic level throughout the outcrop area, the level is considered a fairly dependable time marker. For this reason the top of the Tanyard has been placed at the top of the granular chert zone in the subsurface sections. This probably results in a slight error in correlating with the surface sections, but it does offer a consistent correlation of the top of the formation from well to well.

The criteria for recognizing the base of the Tanyard change with geographic location. The problem seems identical with that encountered at the surface. The most easterly wells that pass through the Ellenburger encounter a dolomite section below

beds bearing typical Tanyard residues (Seaboard No. 1 Dawson in Hamilton County, Pl. III, and McCarthy No. 1 Hedrick in Erath County, Pl. IV). Residues from this lower dolomite section are similar in nature and abundance to those from the Pedernales dolomite member of the Wilberns formation. They are therefore considered Cambrian in age. The lower portion of the Tanyard in the easterly sections contains residues typical of the Threadgill member; that is, a proportionately small amount of residue containing waxy shale flakes and some granulated chert. Sections penetrated so far indicate that the Threadgill and Staendebach members of the Tanyard are recognizable in an area extending from western Eastland County to eastern Erath County in width (Pl. IV) and reaching northward at least through the latitude of the north edge of Erath County, The top of the Cambrian in the eastern sections in this area is placed at the point of sudden increase in amount of residue, which contains an abundance of granular, drusy, dolomoldic chert and drusy quartz,

In moving west across the above described area the Upper Cambrian residues change from predominantly chert to silty, argillaceous material. On the basis of evidence seen in the surface sections it is assumed that the change to silty argillaceous residues continues upward into the Tanyard formation in a westerly direction. In eastern Taylor County (Pl. IV) the lowermost Tanyard beds apparently contain such residues. Therefore, from about the longitude of eastern Taylor County westward the Tanyard-Cambrian boundary is best determined on the granularity of the limestone, if present, and abundance of glauconite in the Cambrian, as is the case in the western portion of the outcrop area.

Determination of the Gorman-Honcycut boundary can be made with less assurance than the other stratigraphic boundaries in the section. The sandy character of the Gorman residues is a general guide throughout most of the north Texás area. Accordingly the top of the Gorman is placed at the top of the persistently sandy residue section found in wells in the area extending from the outcrop of the Ellenburger northward to approximately the Archer-Young County line. From about this latitude northward the entire upper

Ellenburger section furnishes sandy residues, and the Gorman and Honeycut have not been distinguished from each other in sections in Archer and Wichita counties.

It is the opinion of the writer that complete Ellenburger sections in the Archer-Wichita County latitude include beds that are younger than Ellenburger at the outcrop. A careful comparison should be made of sample criteria from the subsurface sections with sample criteria from surface sections in the Wichita and Arbuckle Mountains of Oklahoma. Such a comparison may reveal evidence for correlating the extreme north Texas Ellenburger with formations of the Arbuckle group. Should this be so, the youngest Ellenburger beds of north Texas may correlate with the upper Kindblade or lower West Spring Creek formation of Oklahoma. Cloud and Barnes (1948, p. 65) considered the West Spring Creek and approximately the upper third of the Kindblade to be younger than Honeycut at the outcrop.

Plate IV presents the thinning of the Ellenburger group westward by erosional loss at the top and probably depositional thinning of the Tanyard at the base. In Taylor County the Honeycut is missing, 185 feet of Gorman remains, and the Tanyard is 370 feet thick. In Fisher County all the Honeycut and Gorman section is missing, and only 300 feet of Tanyard is present to represent the group.

Prolific oil production from the Ellenburger has been found in only one field to date in north Texas. The K. M. A.-Ellenburger field of Wichita and Archer counties is producing in part from the top of the Tanyard and in part from the base of the Gorman formation (Shell No. 18E Preston, Pl. III). Crowley (1946)² has shown that there is a remarkable difference in the initial potential production from the two formations. The wells completed in the Tanyard in general have initial potentials greater than 1,000 barrels of oil per day. The wells completed in the Corman in general have potentials less than 500 barrels per day.

CENTRAL TEXAS TO WEST TEXAS

Wells drilled in an area extending westward from the southernmost edge of the

²Contact of Crowley's Units B and C is the Gorman-Tanyard contact of this paper. Ellenburger outcrop area encounter a uniform section of Ellenburger beds. Plate V presents the logs and a correlation chart for a traverse extending from Blanco County westward to Crane County. The sections have been correlated almost entirely on the basis of the residue logs. Residues are sufficiently similar to those of the outcrop samples to permit a reasonable correlation.

The Tanyard formation is recognized by the granularity of its chert residues. Most of the Tanyard sections contain a granulated chert zone, similar to the Tanyard of the Llano River section in Mason County (Pl. II). The major change in residue characteristics is the presence of sand in the west Texas sections. The Tanyard is underlain in some sections by a highly glauconitic limestone which is tentatively correlated with the Wilberns formation of central Texas. This correlation admittedly is debatable, it being possible to argue that the glauconitic limestone facies of the Wilberns has simply migrated up the section into beds of Ordovician age. The Tanyard of west Texas is less than half as thick as that at the outcrop. Most Ellenburger sections in west Texas have about 300 feet of Tanyard. Lithologically the Tanyard of west Texas is a medium to coarse dolomite. Good porosity is evident in some portion of most sections.

The Gorman-Tanyard boundary is placed at the top of the section dominated by granular chert as in north Texas. The Gorman residues of west Texas contain a slightly larger percent of granular chert than do those of central Texas, but nongranular chert is a persistent and usually dominant constituent of the Gorman residues. Also the granular chert in the Corman is extremely fine to subgranular. Sand is commonly present, usually as free grains. But a non-granular, sandy chert zone in the upper part of the formation is characteristic. The upper boundary of the Gorman is not a well-defined horizon but must be determined by correlating horizons above and below the boundary, as noted below.

The upper part of the Corman is composed of very fine dolomite, showing little evidence of porosity. The lower portion of the formation is made up largely of medium to coarse dolomite. These lower beds exhibit sporadic zones of porosity. All the Ellenburger above the Gorman in west Texas is here placed in the Honeycut formation. Correlation of the residue logs and thicknesses of sections indicate that only a minor portion of the Ellenburger of west Texas is younger than the youngest Honeycut exposed at the outcrop (Pl. V). There is no marked physical evidence for defining a formation younger than Honeycut. Therefore, it is considered logical to place the youngest Ellenburger beds in the subsurface of west Texas in the Honeycut, even though some of those beds may be younger than any Honeycut at its type locality.

The residues of the Honeycut are largely non-granular chert, with some sand and granular chert. Both sand and granular chert occur in slightly smaller percentages than in the Gorman. One granulated chert zone occurs rather persistently in the lower' part of the formation. This granulated zone and the sandy chert zone in the upper part of the Gorman serve as correlative aids in determining the Gorman-Honeycut boundary. The boundary is placed between the above described zones at a point on the residue logs where there is a slight increase in residue percentage, with an increase in the percentage of granular chert in the residue. This point is usually about midway between the granulated chert zone of the Honeycut and the sandy chert zone of the Gorman.

The Honeycut in most localities contains a sublithographic, white limestone member in its upper part, with the remainder of the formation being composed of very fine to extremely fine dolomite. Any porosity present is correspondingly fine. The formation is rendered more porous and permeable by fracturing in areas of sharp structural disturbance.

The series of sections and logs presented on Plate V are designed to establish a reference cross section to which other sections in west Texas can be referred. The traverse of the sections follows the trend of what may be termed the "normal" facies of the Ellenburger. That is, all the formations are present and contain lithology and residues similar to the type and standard sections of the Ellenburger. Farther westward this facies must grade into the impure dolomitic rocks making up the El Paso formation exposed on

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Beach Mountain in the Van Horn region. Southward the Ellenburger of west Texas must eventually grade largely into the impure limestone facies represented by the Marathon formation exposed in the Marathon uplift. But until the Ellenburger loses its identifying characteristics in adjacent facies it should be possible to recognize its formations by comparison of insoluble residue logs.

Excellent oil production has been obtained from all portions of the Ellenburger section in west Texas. Both fracture and interstitial porosity occur, depending on the geological history of the structure. The sections on Plate V include the Humble No. 4-C Fee, a well in the Todd field of Crockett County where Ellenburger production is from the Honeycut formation.

SYSTEMATIC DISCUSSION OF RESIDUES³

CHERT

The term "chert" is applied to the cryptocrystalline varieties of quartz found in the residues, regardless of color. The chert is composed mainly of microscopic fibers of chalcedony or particles of quartz, or a mixture of both. The quality of the chert that has proven to be diagnostic in studying the Ellenburger is texture. Three main textural groups found in the Ellenburger residues have been designated as smooth, granular, and chalky. The majority of the chert encountered can be readily placed in one of the major groups, with a small minority showing intermediate nondeterminate characteristics. Each of the textural groups shows variations in appearance, making it possible to subdivide each group into recognizable types.

The chert described as smooth textured has a conchoidal to flat fracture surface devoid of roughness. It has no distinctive structure, crystallinity, or granularity. Some of it is botryoidal in form. The types recognized are ordinary, chalcedonic, and porcelaneous. The ordinary smooth chert typically has a flat fracture, may be any color but is usually white or gray, is mostly opaque but some variations show slight translucency. It occasionally presents a mottled or flocculated appearance. It is generally homogeneous but may have slight evidence of granularity or crystallinity. The chalcedonic smooth chert is transparent to translucent and is described as smoky, milky, waxy, or greasy in appearance. Variations are tinted with various colors, but the most common is light blue gray. It can also be finely mottled or have flocculated inclusions. The fracture is more usually conchoidal than flat." The porcelaneous type of smooth chert has a very smooth fracture surface, is opaque to subtranslucent, and is typically chinawhite, resembling china-ware or glazed porcelain. It grades into an extremely fine granular chert resembling unglazed porcelain.

The granular chert found in the Ellenburger residues is composed of distinguishable grains, granules, or druses. It has an uneven, or rough, fracture surface and a dull to slightly glimmering luster. Fragments of the chert range from hard to soft. Two types recognized are designated as compact and granulated. The compact type is mostly homogenous, has relatively uniform-sized grains, and some of it ap-pears saccharoidal. The grain size varies from an extremely fine type that is gradational from the smooth porcelaneous up to a size that is easily distinguishable under 10-power magnification. The grains rarely reach 0.5 mm in size. The granulated type of granular chert is composed of tiny grains or granules of chert tightly to loosely held together in small, irregular masses or fragments. The individual grains making up the small masses may appear saccharoidal, grade from angular to drusy, and range from extremely fine up to grains 0.5 mm in size. This type of chert seems always to be well disseminated through the enclosing carbonate rock and is found more commonly in dolomite than in limestone.

The chalky chert in the Ellenburger residues is so called because of its textural resemblance to chalk. Much of it bears a strong resemblance to tripolite and may be tripolitic in origin in that it may represent an aggregation of siliceous particles, the particles being possibly colloidal in size. Some of the chert may be an alteration product formed by the leaching of other

³In June 1946 a group of geologists familiar with residue work mot in conference in Midland, Texas, and worked out a systematic classification based on appearance and agreed on descriptions for materials found in residues. All descriptions of the appearance of Ellenburger residues follow the outline published as a result of the conference. See Ireland and others (1947).

kinds of chert, notably the smooth varieties. Fragments have an uneven or rough fracture and usually a dull or earthy appearance. Some varieties are extremely fine porous.

All the Ellenburger cherts exhibit a variety of structures. The most common structural features are described as porous, dolomoldic, oolitic, pseudo-oolitic, oomoldic, lacy, drusy, banded, and dolomorphic. In addition, the chert may be sandy, silty, spicular, quartzose, and very rarely fossilferous.

Dolomoldic chert contains dolomolds, a dolomold being a rhombic opening left in an insoluble residue by the solution of an embedded dolomite crystal (Ireland and others, 1947). The occurrence of the dolomolds may vary from widely scattered individual openings to very closely spaced, interconnecting openings. The latter occurrence results in a skeleton-like structure and is described as siliceous skeletal dolomolds. The skeletal dolomolds are interstitial chert filling what otherwise would be pore space between the dolomite crystals.

Oolitic chert contains ooliths (DeFord and Waldschmidt, 1946). Ooliths as applied to insoluble residues of the Ellenburger are spheroidal bodies with a nucleus or central mass enclosed by one or more surrounding layers of the same or different material. The ooliths are typically siliceous but are very rarely composed of clay, pyrite, or limonite. The ooliths in some cases are of different color and texture than the surrounding matrix. The occurrence varies from widely scattered to very closely spaced, the latter occurrence forming essentially a siliceous oolite. The ooliths also occur free or in loosely held aggregates. The ooliths grade from very small to large but generally average about 1 mm in size. Internally the ooliths have either a distinct or indistinct nucleus surrounded by concentric, radiate, or massive material. Externally the ooliths are smooth or covered with very fine druse.

Pseudo-oolitic chert contains rounded, pellet-like bodies with no peripheral layer, some of which do not exhibit sharp distinction between pellet and matrix. Origin of this particular structure is obscure and possibly varied. The pellets possibly are ooliths with the outside layer absorbed in the process of replacement, or they may result from replacement of pellet limestone. A similar structure has been described as pseudospicular because of its resemblance to the pattern formed by the spicular meshwork of certain lithistid sponges (Cloud, Barnes, and Bridge, 1945, p. 136).

Both colitic and pseudo-colitic chert may be largely the result of replacement. The replacement possibly involved the spherical bodies only, or it could have affected both bodies and matrix.

Oomoldic chert contains spherical cavities formerly occupied by ooliths, each opening being an oomold. The size of the oomolds ranges from microscopic to easily megascopic, and the average is approximately 1 mm in diameter. Their occurrence varies from widely scattered to very abundant. If the oomolds are so abundant that the constituent chert makes up 25 percent or less of the mass, the chert is described as skeletal oomoldic.

Lacy structured chert is a highly porous chert with irregular openings, the constituent chert comprising 25 percent or less of the mass. This chert is largely interstitial material, filling irregular spaces in the carbonate rock that would otherwise be porous. Rarely, the lacy chert appears to be organic in origin, possibly associated with some form of plant life. In the latter case the chert is not interstitial in character but simply embedded.

Drusy chert is in part, at least, encrusted with subhedral quartz crystals. Size of the crystals ranges from microscopic to megascopic. The granular cherts are much more commonly drusy than the smooth or chalky cherts.

Banded chert is made up of layers with either textural or color variations. The layers are microscopic or megascopic. They are flat, as in banded vein filling, or curved, suggesting concretionary structure.

Dolomorphic chert is formed from previously dolomitic or dolomoldic chert whose dolomite rhombs have been replaced or their molds refilled with another mineral. "Dolomorphic" is synonymous with "dolocastic" as used by Cloud, Barnes, and Bridge (1945) but is preferable in order to avoid confusion with "dolocastic" as used by other writers (McQueen, 1931). To be present in insoluble residues, dolomorphic chert must have the dolomite rhombs replaced or their molds filled with an insoluble mineral. It is rare in Ellenburger residues.

Spicular chert contains siliceous or silicified spicules, chiefly sponge spicules. Generally the spicules occur as individual needle-like bodies embedded in the chert, but in some samples the meshwork formed by a mass of the spicules is preserved. Occurrence of the spicular chert is indicative but not absolutely diagnostic of certain levels of the Ellenburger.

Quartzose chert contains inclusions and veins of crystalline quartz, ranging from microscopic to megascopic.

QUARTZ

Quartz in the Ellenburger insolubles means the nondetrital occurrence of the crystalline, colorless, clear to cloudy mineral. It occurs in euhedral, subhedral, and anhedral forms. The euhedral quartz occurs as doubly terminated crystals embedded in the carbonate rock and as free crystals in the residues. The crystals vary from microscopic to megascopic in size and from fine, needle-like to short and stubby in shape. The subhedral quartz occurs chiefly as drusy encrustations on rock surfaces. When the encrustations have been removed from carbonate rocks by solution they occur in the residue as free or clustered, partially developed crystals. The anhedral quartz shows no outward crystal development. It exhibits textures ranging from smooth, glassy to granular and granulated.

CLASTICS

The constituents of clastic origin found in the Ellenburger residues are the argillaceous materials—clay and shale; silty material; and sandy material. These constituents occur generally disseminated through the carbonate rock but may occur as small, discontinuous seams or partings. All the material included in this category is considered to be detrital in origin. This distinction becomes difficult, and many times a matter of individual judgment when dealing with finely divided, probably precipitated silica and fine silt.

A large portion of the argillaceous material in Ellenburger residues is clay-like in appearance and is here referred to as clay. It occurs interstitially in limestone and dolomite, in two principal forms. One form is best described by the term spongy and occurs in porous, earthy, fragile masses. The other form is flaky, the flakes being usually very crinkled and irregular in shape. The clay exhibits many of the structural forms found in chert including dolomoldic, skeletal dolomoldic, lacy, oolitic, or comoldic forms. It also contains silt or sand grains in some occurrences. The color generally is dull gray to white, but shades of purple, red, and green occur. The flaky form is much more abundant in limestone than in dolomite.

The shale in Ellenburger insolubles occurs as thin partings and interstitial fillings in the limestone and dolomite. In texture it grades from waxy to granular and is hard to soft in character. The majority of the shale is some shade of gray, but green is common, with maroon, brown, and purple being more rare. Laminated, porous, dolomoldic, oolitic, and oomoldic forms occur. Impurities make the shale silty, sandy, micaceous, pyritic, or glauconitic.

Silty material in the residues is characterized by detrital quartz grains of silt size, that is, larger than clay particles, and with a maximum of approximately 1/16 mm. The silt occurs in the residue as loose grains or as poorly to well consolidated aggregates. The aggregates are usually held together by an argillaceous matrix. In some occurrences the silty material is sandy, glauconitic, pyritic, or micaceous. In addition the well-consolidated aggregates exhibit porous, dolomoldic, oolitic, or oomoldic structure.

The sandy portion of the clastic material in the residues occurs largely as loose grains in the residues but also is present in poorly to well-consolidated aggregates. The aggregates are in a shale, chert, or quartz matrix. The grains are, with rare exception, composed of clear quartz. They range in size from just above silt size to more than 2 mm. The most common form is subrounded to rounded with a frosted surface. A minor portion of the sand found in the Ellenburger is angular. Still fewer of the grains exhibit secondary growth. The sand occurs in laterally discontinuous zones in the carbonate rocks. The amount of sand in the zones varies from widely scattered grains to very abundant. But rarely does the sand become abundant enough to form sandstone. Some zones are gradational at the top and bottom, while others show an abrupt disappearance of sand from the section at either the top or bottom of a zone.

ACCESSORY MATERIALS

Accessory materials are those that occur sporadically, or more or less continuously in small amounts, in the Ellenburger residues. These materials are both mineral and organic in nature. Brief mention of each follows.

Pyrite is the commonest accessory. It occurs as interstitial filling and as individual crystals. The crystals range from microscopic to megascopic in size. The pyrite is often associated with waxy shale partings in the rock, occurring as minute crystals embedded in the shale. Some of the mineral called pyrite may actually be *marcasite*, as distinction between the minerals in small occurrences is difficult.

Mica, commonly muscovite, more rarely biotite, occurs as very small, thin flakes. The flakes occur either free or embedded in the shale or any of the aggregates found in the residues.

Feldspar appears rarely in the residues as minute, light-colored crystals, probably sedimentary in origin. Feldspar also occurs as detrital grains at certain places in the Ellenburger.

Limonite is more common in residues from surface samples than from well samples. Commonly it appears to be replacements of pyrite,

Anhydrite appears erratically in subsurface sections of the Ellenburger, usually in a clear, crystalline form. The common occurrence of anhydrite in drilling muds makes it difficult to determine small amounts in place in well samples.

Glauconite occurs in the residues usually as small grains, either free or embedded in shale or silty aggregates. It occurs also as interstitial filling and is dolomoldic when thus occurring in dolomite. Glauconite is more common in the lower portion of the Ellenburger than elsewhere. *Millerite*, in typical needle-like crystal clusters, occurs very rarely.

Fossil fragments are too rare and nondiagnostic to offer faunal evidence with correlative significance. The commonest fragments are sponge spicules, either free or embedded in chort. Both simple, needlelike forms and six-pointed, hexactinellid types are found. Rarely small gastropods and fragments of brachiopods in siliceous replacements are found. As noted above, a few conodonts have been obtained from acetic acid residues of limestone samples.

Beekite is a form of chalcedonic chert associated with fossil replacement. The chert forms a series of rough, buttonshaped bodies made up of roughly concentric rings of chert, partly or wholly covering the fossil surface. The material is frequently observed on the outcrop, and fragments of the beekite bodies appear sporadically in the residues.

Other insoluble minerals and materials than those mentioned above undoubtedly occur in trace amounts in the Ellenburger residues. An exhaustive study and identification of all the more rare constituents of the residues has not been attempted by the writer. Goldich and Parmelce (1947) have published detailed analyses of samples from portions of Ellenburger surface sections.

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SAMPLE DESCRIPTIONS

GORMAN FALLS SECTION, SAN SABA AND LAMPASAS COUNTIES, TEXAS

Honeycut formation, 0-325 feet—	Fee low	-
Dolomite—extremely fine, very light gray	۵.	5
RESIDUR: Clay-sponge-like, yellow. Silt-fine quartz. Chert-trace smooth, white, grading to chalky. Sand-few fine, rounded frosted grains. Spicule-fine, needle-like.	0-	J
Dolomite—extremely fine, ash-gray, 50 percent. Limestone- smooth ash-gray, 50 percent RESIDUE: Clay—sponge-like and flaky, white. Silt—quartz, grading up to very fine sand. Sand—small amount, grading to silt. Chert—trace white, smooth, very roughly nodular.		10
Limestone—smooth, very light gray RESIDUE: Clay—crinkled llaky, light buff, some white. Chert—trace smooth, light buff. Silt—trace,	<u>1</u> 0–	15
Dolomite—extremely fine, ash-gray RESIDUE: Clay—sponge-like, lacy.		20
Limestone—smooth, dove-gray, 90 percent. Dolomite—extremely fine, light gray, 10 percent RESIDUE: Chert-ordinary smooth and porcelaneous, white; chalky, finely porous;	20-	25
some fine granular, white. Spicules—abundant, needle-like, curving, and finely branch- ing. Clay—sponge-like, white to buff.		
Dolmite—extremely fine, light cream-colored, 50 percent. Limestone—smooth, gray, 50 percent RESTUDE: Clay—sponge-like and flaky, white, some maroon. Chert—small amount	25–	30
hard, chalky. Spicules—fine, needle-like, some with extremely fine end filament. Dolomite—extremely fine, light cream and ash-gray		
RESIDUE: Clay—sponge-like, grayish, lacy. Chert—small amount chalcedonic, con- cretionary. Quartz—trace anhedral, rough fragment. Silt—small amount, very fine. Spicules—few fine, needle-like.	-06	99
Dolomite—extremely fine, light gray. Limestone—some smooth, gray. Chert—some smooth, gray to brown RESIDUE: Chert—smooth, brownish gray, slightly chalcedonic, with finely flocculent structure, in part pseudo-oolitic. Silt—small amount very fine. Clay—small amount gray, sponge-like. Spicules—few fine, needle-like.	35–	40
Limestonesmooth, gray	40-	45 ⁻
Limestone—smooth, dove-gray. Chert—small amount grayish white RESIDUE: Chert—smooth, buff; fine granular, grayish, in part quartzose. Clay— crinkled flaky, brown. Spicules—few very fine, needlo-like.	45 -	50
Dolomite-extremely fine, light gray, 80 percent. Limestone-smooth, gray, 20 percent RESIDUE: Clay-sponge-like, lacy, and crinkled flaky, gray. Spicules-fine, needle-like.	50-	55
Dolomite—extremely fine, light gray, 90 percent. Limestone—smooth, light gray, 10 percent RESIDUE: Clay-small amount crinkled flaky, white, buff, and gray. Chert-trace chalcedonic.		
Limestone—smooth, light gray RESIDUE: Chert fine granular, white, porous, small amount quartzose, trace pseudo- oolitic. Quartzfew fine to very fine angular grains. Clay trace flaky,	60-	65
Dolomite—extremely fine, light huff, 90 percent. Linestone—smooth gray, 10 percent. RESIDUE: Sand-abandant, very fine, angular; fine to coarse, well rounded and frosted. Chert—chalcedonic, drusy in part. Clay—small amount sponge-like, lacy.	65	70
Dolomite—extremely fine, light gray, 50 percent. Limestone—extremely fine to smooth, light gray, 50 percent RESIDE: Chert—chalky, soft, porous. Clay—sponge-like, lacy, and crinkled flaky, white and huff.	70–	75
Limestone—smooth, light gray and gray RESIDUE: Chert—small amount very rough, concretionary, white. Clay—small amount crinkled flaky, buff.	75–	80
Limestone—smooth, very light gray RESTOUE: Clay—crinkled flaky, white, some buff. Chert—small amount very roughly concretionary smooth white. Sniculas—few extremely fine	80–	85

Report of Investigations-No. 11

Feet below top Limestone—smooth, dovo-gray
RESIDUE: Clay—crinkled flaky, light buff to white. Chert—trace smooth, light buff. Limestone—smooth, light, slightly buff-gray
smooth textured, gray.
Dolomite—extremely fine, light cream
Dolomiteextremely fine, very light cream 100-105 RESIDUE: Clay-sponge-like, gray, in part lacy.
RESIDUE: Clay-sponge-like, gray, in part lacy. Limestone-smooth, gray 105-110 RESIDUE: Clay-sponge-like to slightly compact, white; crinkled flaky, buff. Silt- small amount.
Dolomiteextremely fine, very light gray, 50 percent. Limestonesmooth, gray, 50 percent
fine granular, grayish, in roughly botryoidal fragments. Limestonesmooth, gray, 90 percent. Dolomiteextremely fine, very light cream, 10
percent
Limestone-smooth, gray, 60 percent. Dolomite-extremely fine, very light cream, 20 percent. Chert-smooth, gray, 20 percent
RESIDUE: Chert-chalcedonic, gray and maroon, in part spicular; small amount rough, fine granular, gray. Silt-small amount. Clay-small amount crinkled flaky. Spicules - few needle-like.
Limestone -smooth, gray 125-130 RESIDUE: Clay—crinkled flaky, brown and maroon. Chert—small, granular fragments, apparently fragments of silicified fossils and small concretions. Spicules—abundant needle- like, few with finely branching ends.
Limestone-smooth, gray 130-135 RESIDUE: Chert-smooth, in part slightly chalcedonic, in part porcelaneous, traces psuedo-oolitic structure, some spicular, white to buff. Clay-flaky and sponge-like, light buff. Spicules-fine, needle-like.
Limestone—smooth, gray, in part with uneven, fragmental-appearing structure. Dolomite —trace very fine, light gray135-140 RESIDUE: Chert—smooth, light buff to white, spicular; chalky, hard; some granular, white. Clay—sponge-like, some flaky, white, huff and brown in part dolomoldic. Spicules
fine, needle-like. Limestonesmooth, gray, in part marcon splotched, in part with indistinct, uneven
structure 140-145 RESIDUE: Chert-smooth, variegated powder-white and grayish chalcedonic, spicular in part. Clay-sponge-like, maroon and buff. Spicules-few, needle-like.
Limestone—smooth, gray
Limestone—smooth, gray
Limestone—smooth, gray
Limestone-smooth, gray, in part with indistinct fragmental structure
Limestonc—smooth, gray
Limestone—smooth, light gray, some pivkish maroon

Correlation Between Sections of Ellenburger Group

Feet Lalase ter
below top
Limestone—smooth, dove-gray
Limestone—smooth, dove-gray. Dolomite—small amount very fine, light gray
Limestone—smooth, gray, in part finely dolomitic
Dolomite—very fine, light gray. Limestonc—some smooth, gray, dolomitic
Dolomite—extremely fine, light buff-gray, 80 percent. Limestone—smooth, very light gray, 20 percent
Quartz—anhedral, rough to nodular, and subhedral, drusy: Spicules—fine, needle-like.
Limestone—smooth, very light gray, 90 percent. Dolomite—extremely fine, light buff, 10 percent
anhedral. Limestone—smooth, light gray, 90 percent. Chert—smooth, light tan, 10 percent205-210 RESIDUE: Chert -smooth, white and tan; fine granular, grayish, and finely laminated,
tan. 910 915
Limestonc—smooth, dove-gray210-215 RESIDUE: Clay-wery crinkled flaky, tan to white. Chert—trace chalcedonic. Spicules —trace fine, needle-like.
Limestone—smooth, very light gray
Linestone—smooth, dove-gray
Limestone—smooth, dove-gray 225-230 RestDUE: Chert—smooth, white to fine granular porcelaneous; some fine granular, slightly translucent-gray, scattered porosity. Spicules—few extremely fine.
Limentary sweeth grow 50 percent Delemite extremely fine light slightly huff-gray
50 percent
Limestone—smooth, dove-gray. Chert—some smooth grayish white235-240 RESIDUE: Chert—smooth, grayish white, and chalcedonic. Quartz—rough, granular, some lacy. Clay—flaky, greenish, buff, and white. Silt—small amount very line. Spicules
—fine, needle-like.
Limestone—smooth, light gray. Chert—some smooth white
Limestone-smooth, light gray, 80 percent. Dolomite-extremely finc, light gray, 20 percent
small amount.
Limestone—smooth, light gray. Dolomite—small amount finc, light gray
Limestone—smooth, light gray, trace finely dolomitic
Limestone—smooth, light gray, some finely dolomitic 260-265 RESIDUE: Clay—flaky, light greenish. Silt—coarse to very fine.
Dolomite—extremely fine to very fine, light buff and light pinkish gray. Limestone—trace light gray, dolomitic
RESIDE: Quartz—subhedral, in irregular, porous masses. Chert—chalky, soft, white. Sand—fine to very fine, lightly frosted. Silt—very fine. Clay—white, finely fragmental dolomoldic.

Feet

below to
Dolomite—very fine, light gray
Quartz-subhedral, in rough, irregular fragments.
Limestonesmooth, light gray
RESIDUE: Clay—crinkled flaky, white, some pink. Sand—fine to medium coarse, well rounded, lightly frosted. Chert—small amount smooth, white.
Dolomite—extremely fine, buff, some light gray
Dolomite—very fine to fine, light grây, in part with scattered sand grains
Limestone—smooth, light gray. Chert—trace smooth, grayish white and brown
Limestone—smooth, dove-gray, 90 percent. Dolomite—very fine, buff, 10 percent
Dolomiteextremely fine to very fine, brown to light gray
Dolomite—extremely fine, light buff-gray, 90 percent. Chert—smooth, white, oolitic, 10 percent
grains; fine granular, translucent-grayish, extremely fine dolomoldic. Sand—few coarse, rounded, frosted grains.
Dolomite—very fine, gray. Limestonesmall amount, smooth, very light gray
Limestone—smooth, grayish white, 50 percent. Dolomite—very fine, gray, 50 percent315-32 RESIDUR; Clay—bulf-brown, finely skeletal dolomoldic, and very finely fragmental. Chert—trace smooth white.
Dolomite—very fine, buff-gray, in part highly arenaceous, 70 percent. Limestone—smooth, gray to maroon mottled, in part arenaceous, 30 percent, 320-32 RESIDUE: Sand—abundant, coarse to very fine, well rounded, very finely frosted, some grains lightly colored by maroon fillings in frosting. Clay—small amount light buff, sponge-like, in part lacy.
Gorman formation, 325-745 feet-
Limestone—smooth, light gray, in part dolomitic
Limestonesmooth, dove-gray, with few embedded, frosted sand grains
Limestone—smooth, light gray
Limestone—smooth, very light gray
Limestonc-smooth, gray RESIDUE: Clay-flaky, gray, slightly granular. Quartz-very delicately lacy fragments.
Limestone—smooth, dove-gray
Limestone—smooth, light gray
Limestone smooth, light tan-gray, 80 percent. Dolomite - fine, gray, 20 percent
Limestone—smooth, light gray finely splotched with brown

Correlation Between Sections of Ellenburger Group

Feet below top Dolomite-very fine, light buff-gray____ 370-375 RESIDUR: Chert--smooth, white; trace hard chalky; trace fine granular, white. Clay-buff, very finely fragmental, and brown, skeletal dolomoldic. Dolomite-very fine, light gray, buff, and brown RESIDUE: Clay-buff, extremely fine Iragmental. Chert--trace fine granular, dolo-.....375-380 moldic, drusy. Sand-few fine to medium, rounded lightly frosted grains. Dolomite-very fine to extremely finc, light buff-gray, 90 percent. Chert-smooth, Dolomite-very fine, light gray, 50 percent. Limestone-smooth, very light gray, 50 percent RESIDUE: Clay-flaky, greenish brown. Limestone-smooth, gray RESIDUE: Chert-chalky, hard; smooth, white, trace oolitic. Clay-sponge-like, white, flaky, greenish brown. Silt-small amount.405-410 Limestone-smooth, dove-gray RESIDUE: Clay-flaky, gray and buff. Sand-small amount fine, rounded, very lightly frosted. Chert-trace smooth, white, finely solitic. RESIDUE: Clay-flaky, whitish gray. Quartz-trace drusy. Dolomite-fine to medium fine, gray..... RESIDUE: Clay-flaky, buff to gray. Sand--fine to very fine, rounded, very lightly frosted. Silt--coarse to very fine. Chert-trace smooth, white. Limestone-Smooth gray in part dolomitic, 80 percent. Dolomite-extremely fine, buff.....420-425 RESIDUE: Clay-flaky white, and extremely fine fragmental, brown. Limestone-smooth, gray RESIDUE: Chert-smooth, white, in part with very finely flocculent structure. Sand--small amount fine, rounded, frosted. Sand-few fine, slightly frosted grains. Limestone-smooth, gray, in part indistinct, finely fragmental structure..... 435-440 RESIDUE: Sand-very fine to medium, rounded, frosted. Silt-coarse to very line. Clay-flaky, brown to buff. Limestone-smooth, light gray, 50 percent. Dolomitc-very fine, light gray, 50 percent....440-445 RESIDUE: Clay-small amount, extremely fine, fragmental. Dolomite-very fine, light gray .445-450 RESIDUE: Chert-smooth to slightly chalky, very finely scattered to abundant dolomoldic; some very fine granular white, very finely dolomoldic. Limestone—smooth, gray RESIDUE: Clay—flaky, white. Shale—small amount dark green, waxy. Sand—small 450-455 amount fine, rounded, lightly frosted. Silt-coarse. Limestone-smooth, gray, 70 percent. Dolomite-fine, gray 20 percent. Chert-smooth, gray, oolitic, 10 percent..... donic, highly oolitic and partly arenaceous; fine granular, very finely porous, white. Sand-small amount fine, rounded, frosted. RESIDUE: Clay-crinkled flaky, greenish gray and white. Chert-chalky, soft, grayish white. . 465-470 Dolomite---extremely fine, light gray...... RESIDUE: Clay-small amount sponge-like, gray. Chert-small amount smooth, white. RESIDUE: Silt-fine. Sand-few fine, frosted grains. Dolomite-extremely fine, light gray, some buff RESIDUE: Silt-fine to coarse, angular, loose and in soft, argillaceous aggregates. Sand -very fine, rounded, very slightly frosted.

27

	Feet
Limestone-smooth, very light gray. Dolomite-small amount very fine, buff	<i>below top</i> 480–485
RESIDUE: Chert—smooth, gray to white, in part fine to medium, colitic, part pseudoolitic, contains scattered sand grains, has scattered porosity. Sand—very fine to mediu rounded, very slightly frosted. Clay—crinkled, greenish white.	ło-
RESIDUE: Clay—crinkled flaky, white, gray, and buff. Sand—fine to very fine, we rounded, very slightly frosted. Silt—small amount. Spicales—fine, needle-like. Shale	485-490 ell
trace gray, slightly waxy. Limestone—smooth, dove gray	490-495
RESIDUE: Clay—crinkled flaky, buff to gray. Sand—very fine, slightly frosted.	. 170 170
RESIDUE: Clay-crinkled flaky, white, some greenish gray, some muddy gray. Sand very fine, very slightly frosted.	
Limestone—smooth, light gray	500–505 all
Limestone—smooth, light gray RESIDUE: Clay—crinkled flaky, slightly greenish buff, brown, some white. Silt—ve small amount.	505–510 ry
Limestone—smooth, very light gray	510-515
RESIDUE: Chert—chalky, white, porous; smooth, white. Clay—flaky, greenish grasome white. Sand—small amount very fine, slightly frosted.	ıy,
Limestone—smooth, dove gray RESIDUE: Clay—crinkled flaky, white, greenish gray, buff. Silt—small amount coars Sand—very small amount very fine, very slightly frosted.	515–520 se,
Limestone—smooth, very light gray. RESIDUE: Clay—crinkled flaky, light gray, greenish gray, and buff, in part slight silty. Silt—small amount very fine.	520–525 Ily
Limestone—smooth, very light gray	525-530
RESIDUE: Clay—crinkled flaky, white to slightly greenish. Limestone—smooth, dove-gray, in part arenaceous, 90 percent. Chert—smooth, light bu 10 percent	ff, 530–535
RESIDUE: Chert-smooth, white to very light buff; chalky, soft, white. Sand-coar to fine, rounded, frosted. Clay-flaky, greenish gray and white, very finely fragmenta	se
Limestone—smooth, light grav	535 - 540
RESIDUE: Clay-flaky, white to buff. Sand-well sorted, fine, rounded, lightly froste Limestone-smooth, light gray	ed. 540–545
RESIDUE: Clay-flaky, slightly greenish, light gray, bpff. Sand-very fine, slight frosted, grading to coarse silt. Silt-small amount coarse.	ly
Limestone-smooth, very light gray	
RESIDUE: Chert-gradational between chalky and fine granular, hard, white, fine porous; fine granular, translucent-gray, in part porous, and in part pseudo-oolitic.	ly
Limestone—smooth, light gray, in part indistinct fragmental structure RESIDUE: Clay—flaky and sponge-like, greenish buff and white.	
Limestone—smooth, very light gray	555-560
chalky. Silt-very fine.	
Limestone—smooth, light gray, some maroon staining. RESIDUE: Clay—flaky, greenish buff and salmon-pink. Silt—small amount very fin Chert—trace smoky, chalcedonic.	560–565 1e.
Limestone—smooth, light gray	565–570
Limestone—smooth, light gray	570–575
Limestone—smooth light gray, with few embedded sand grains RESIDUE: Sand—coarse to fine, rounded to subrounded, lightly frosted. Clay—crinkle flaky, brown to buff.	575–580 ed
Limestone-smooth, light gray. Chert-small amount white, calcarcous.	580585
RESIDUR: Chert—chalky, white, finely porous, some gradational to fine granular; fu granular white; smooth, white. Clay—crinkled flaky, greenish gray to buff.	ne
Limestone-smooth, light gray, 50 percent. Dolomite-coarse light gray, 50 percent	. 585–590

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Correlation Between Sections of Ellenburger Group

Feet
below top
Limestone—smooth, light gray
Limestone-smooth, light gray, 90 percent. Chert-smooth, white, 10 percent
Limestone—smooth, very light gray
Limestone—smooth, light gray, 50 percent. Dolomite—extremely fine, light gray, 50 percent
Dolomiteextremely fine, gray
Delomite—extremely fine some fine, light gray, 60 percent, Limestone—smooth, light
gray, 40 percent. Chert-trace, grayish white
Limestonesmooth, light gray. Dolomitesmall amount medium, light gray
Limestone—smooth, dove-gray
Limestone—smooth, light gray
Limestone—smooth, light gray
Dolomite—very fine to extremely fine, light gray
Dolomiteextremely fine, very light gray645-650 RESERVE: Sand-very fine, slightly rounded and frosted to extremely fine, angular.
Dolomite—extremely fine to very fine, light gray, some buff
Dolomite—extremely fine, buff; and fine, light gray
Limestone—very smooth, white, 80.percent. Chert—smooth and chalky white, 20 percent660-665 RESIDUE: Chert—chalky, hard to very soft; smooth, white.
Dolomite—extremely fine, light buff-gray; and medium fine, gray
Dolomite—extremely fine to fine, light buff-gray
Limestonesmooth, light gray, 50 percent. Dolomiteextremely fine to fine, buff-gray, 50 percent
Limestonc—smooth, light gray, in part arenaccous. Chertsome smooth, gray
Limestone—smooth, light gray

29

Feet
Limestene energie licht man
Limestone—smooth, very light gray
Limestonc-smooth, light gray
Limestone—smooth, gray, some maroon, 80 percent. Chert—smooth, gray, oolitic, 20 percent
RESIDUE: Chert—smooth, gray, oolitic, few embedded sand grains; smooth, white. Clay —sponge-like and flaky, greenish white and brown. Sand—small amount fine, rounded, frosted. Silt—small amount coarse to fine.
Limestone—smooth, light gray, in part with indistinct fragmental structure, 60 percent. Dolomite—extremely fine, light gray, 40 percent
Dolomite—extremely fine, grayish white710-715 RESIDUR: Silt—fine, abundant. Chert—chalky, white. Clay—small amount very deli- cately sponge-like.
Dolomite—extremely fine, very light buff, and grayish white
Dolomite—fine to extremely fine, light gray, small amount pink
Dolomite—very fine to extremely fine, very light gray
Dolomite—extremely fine to very fine, grayish white
Dolomite—extremely fine to very fine, light gray, in part sandy
Limestone—smooth, light gray
Spicewood Creek Section, San Saba County, Texas
Forman formation, 0–57 feet—
Dolomite—extremely fine to very fine, light, very slightly buff, 60 percent. Limestone—extremely fine, grayish white, 30 percent. 0-5 RESIDUE: Silt—coarse to fine. Clay—in very fine particles. Sand—few very fine grains. 0-5 Chert—trace smooth, white, in part chalcedonic. 0-10
Dolomiteextremely fine, light, very slightly buff; and very fine, very light gray
Dolomite—extremely fine, light buff; and very fine, very light gray with scattered pinkish stains. Chert—trace ordinary to porcelaneous smooth
Dolomite—extremely fine, very light gray, with irregular, light buff coloration 15- 20 RESIDUE: Chert—chalky, argillaceous, extremely fine, abundant to skeletal dolomoldic. Clay—sponge-like, extremely fine skeletal dolomoldic. Sand—very fine, rounded, frosted, loose, and in small aggregates. Silt—Joose.
Dolomite—very fine, very light gray. Chert—small amount smooth white, in part finely dolomitic; trace extremely fine granular chert

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	Feet below top
Dolomite—extremely fine to very fine, very light gray RESIDUE: Claysponge-like and flaky, finely dolomoldic. Sand- very fine, ligh frosted. Silt—loose.	25 30
Dolomite—extremely fine, unevenly light buff colored, 50 percent; very fine to fi light gray, 50 percent RESIDUC: Chert—smooth, white, quartzose and drusy, scattered to abundant fin dolomoldic; chalky, white, scattered dolomoldic. Quartz -subhedral, finely drusy. C —sponge-like, extremely fine to very fine skeletal dolomoldic. Silt—loose.	30–35 cly
 Dolomite—extremely fine, grayish white, in part with embedded, very fine to f sand grains RESIDUE: Sand—unassorted, very fine to coarse, rounded, frosted. Clay—extrem fine fragmental skeletal dolomoldic. Silt—loose. 	35- 40
Dolomite—extremely fine, light gray. Chert—small amount smooth, gray, in part w very fine pseudo-oolitic structure in chalcedonic matrix	40 45 ure
Dolomite—extremely fine, very light gray. RESIDUE: Chert—ordinary smooth, white. Sand—few very fine, lightly frosted grai Silt—small amount loose. Clay—few flakes.	45- 50 ins.
Dolomite—very fine, light buff-gray. Chert—small amount smooth, grayish white a finely banded	50– 55 ng, unt
Tanyard formation, Staendebach member, 57-460 feet— Dolomite—very fine to extremely fine, light gray with uneven, light buff coloration RESTOUR: Chert—smooth, white, very finely abundant dolomoldic; extremely is granular, slightly translucent, finely abundant dolomoldic, in part quartzose and dru Quartz—subbedral, finely drusy. Clay—sponge-like, very finely fragmental skeletal do moldic. Silt—very small amount loose.	ine 1sy. plo-
Limestone—smooth, dove-gray, in part with fine, indistinct pellet structure. Dolomit trace fine, gray RESLDUE: Clay—flaky, white, few fine dolomolds. Sand—few very fine, frosted gray Cherttrace botryoidal chalcedonic.	e— 60- 65 ins.
Limestone—smooth grayish white and dove-gray, 85 percent. Chert—smooth, blu white and trace porcelaneous, 15 percent. RESIDUE: Chert—smooth, chalcedonic to ordinary, white; small amount chalky; sn amount extremely fine granular.	65– 70
Limestone—smooth, dove-gray, in part with indistinct fragmental structure. Chert—sm amount smooth, bluish white, chalcedonic, with few drusy surfaces. RESIDUE: Chert—chalcedonic, quartzose, in part drusy, with trace very finely band Clay—flaky, white and pink stained.	70– 75
Linestonc—extremely fine, dove-gray. Chert—small amount extremely fine granular a chalcedonic	75– 80 ous
Limestone—extremely fine to smooth, dove gray, in part slightly dolomitic, and in p with indistinct fragmental structure. Chert—small amount smooth and some extrem fine granular, white to grayish white	aely 80~ 85 nall red.
Limestone—extremely fine, very light gray, in part siliceous. RESIDUE: Chert—chalky, white, extremely fine porous, soft. Clay—sponge-like a flaky, white.	
Limestone—extremely fine, very light gray, 50 percent. Dolomite—coarse, very light gr in part pink stained, 50 percent. RESIDUE: Chert—small amount finely colitic, chalcedonic; trace smooth, white, Clay finely flaky. Clauconite—trace.	90 95
Linestone—extremely fine, slightly grayish white, in part coarsely dolomitic RESIDUE: Clay—small amount finely fragmental flaky.	95–100

feet
below top
Limestone—extremely fine, very light gray. Chert—small amount smooth, white to gray100-105 RESIDUE: Chert—smooth, white and chalcedonic. Clay—sponge-like and flaky, white.
Linestone—extremely fine to smooth, very light gray. Dolomite—trace medium grained, very light gray
oomoldic. Shale—waxy, light green, in fine fragments.
Limestone—extremely fine, light gray
Limestoneextremely fine to smooth, very light gray RESIDUR: Chertsmooth to extremely fine, white to gray, in sharp contact with lime
Limestone—extremely fine, very light gray, in part containing dolomite rhombs. Dolomite —small amount medium coarse, grayish white. Chert—some extremely fine granular, slightly quartzose, finely colitic
Limestone—extremely fine, very light gray
Limestone—extremely fine, white, in part medium to coarsely dolomitic
Limestoneextremely fine, white, in part medium to coarsely dolomitic
Limestone—extremely fine, very light gray, in part coarsely dolomitie
Limestone—extremely fine to smooth, grayish white
Limestone—extremely fine, grayish white, slightly dolomitic. Chert—small amount ex- tremely fine granular, gray, oolitic, very slightly dolomitic, with the dolomite rhombs and aggregates interrupting the oolitic structure
Dolomite-medium to coarse, grayish white
Dolomite—medium, grayish white, porous, 50 percent. Limestone—extremely fine, grayish white, 50 percent
Dolomite—medium coarse, very light gray. Chertsmall amount, very fine granular, grayish, in part colitic
Dolomite—medium, light buff-gray. Chert—some smooth to extremely fine granular, grayish white, very sparingly dolomitic
Dolomite—coarse to medium, light gray, porous. Chert—small amount variegated smooth and extremely fine granular, white to grayish 175-180 RESIDUE: Chert—smooth, white, quartzose, very sparingly drusy; some variegated smooth to extremely fine granular. Clay—small amount flaky.
Dolomitc—medium coarse, grayish and slightly pinkish white. Limestone—small amount extremely fine, light gray
Limestone—smooth, very light gray, in part intermixed with coarse dolomite. Dolomite— coarse, very light gray, in part intermixed with smooth limestone
Limestone—smooth, very light gray, in part coarsely dolomitic

Correlation Between Sections of Ellenburger Group

Feet below top Limestone-extremely fine, very light gray, with veinlets of pinkish dolomite. Dolomite-some medium to coarse, very light gray_______ RESIDUE: Clay—small amount flaky, white. Chert—trace smooth, white. Dolomite-medium coarse, very light gray with slight rose coloration..... RESIDUE: Clay-small amount fragmental skeletal dolomoldic, some flaky. Chertsmall amount rough, nodular, smooth white. -205 - 210Dolomite-coarse, very light gray..... RESIDUE: Clay-small amount sponge-like, lacy; trace fragmental skeletal dolomoldic. Dolomite-coarse, light gray..... RESIDUE: Clay-small amount sponge-like, flaky. Sand-small amount very fine, subangular, very lightly frosted. Dolomite-medium coarse, very light gray, 80 percent. Limestone-extremely fine, very light gray, 10 percent. Chert-smooth, brown, 10 percent_ 215 - 220RESIDUE: Cheri-smooth, brown, small amount smooth, white. Clay small amount sponge-like, huff-white. Sand-small amount very fine, angular. fine. Dolomite-medium, slightly brownish gray. Chert- small amount smooth, grayish white .. 225-230 -RESIDUE: Chert-ordinary smooth and porcelaneous, white. Spicules-few very fine, needle-like. Dolomite-variegated texture fine to coarse, slightly buff-gray.... RESIDUE: Clay-small amount sponge-like, white. Chert-small amount chalk textured, white; trace chalcedonic. Spicules-numerous very fine, needle-like. Dolomite---medium coarse and medium fine, gray, in part porous. Chert---small amount smooth and porcelaneous, white, in part collitic...... 235-240 RESIDUE: Chert-ordinary white, in part quartzose and drusy; porcelaneous, in part collicie and pseudo-collic. Quartz-small amount subhedral. Ooliths-few free. Spicules -few very fine, needle-like.240-245 Dolomite-medium fine, gray, slightly porous RESIDUE: Clay-small amount 'fragmental, skeletal dolomoldic. Chert-trace chalcedonic, Spicules-few needle-like. Dolomite-medium fine, gray. Chert--small amount smooth, white to gray, in part 245-250 part pseudo-oolitic. Clay-small amount fragmental dolomoldic. 250 - 255Delomite-medium to coarse, gray. Chert-trace smooth, white...... RESIDUE: Chert-smooth, white, in part indistinctly pseudo-oolitic; chalky, creamwhite, scatteringly to abundantly dolomoldic. Clay-sponge-like and fragmental dolomoldic, cream-white. Dolomite -- medium to fine, gray. Chert -- small amount smooth, grayish white with very finely flocculent structure RESIDUE: Chert -- smooth, white to grayish white, in part with poorly developed dolo-255-260 molds, in part finely collitic, some roughly porous and finely drusy. Clay-small amount sponge-like, dolomoldic. Dolomite-coarse, gray, very compact..... RESIDUE: Clay-Sponge-like, highly porous, white and buff. Chert-granular, very slightly translucent and buff, roughly fragmental, sparingly dolomoldic, with some very fine druse; small amount chalky, white porous. 265 - 270Dolomite—coarse, light gray, very compact RESTORE: Chert-granular, gray to buff-gray, in part slightly porons and sparingly dolomoldic. Silt-remarkably clear quartz, possibly chemical in origin. Dolomite- coarse to medium, gray RESIDUE: Chert-chalky, white, in very rough, porous fragments, soft to hard; trace chalcedonic. Clay-sponge-like, white, in very fine, fluffy fragments. Dolomitc-medium coarse, gray, 80 percent. Chert-very fine granular gray, 20 percent 275-280 Resmue: Chert-granular, light, very slightly translucent with some gray, in part slightly porous; small amount drusy. Quartz-very fine, angular grains, possibly chemical in origin. Dolomite-medium coarse, light gray, with scattered pink coloration.... RESIDUE: Chert-fine granular, white, abundant to skeletal dolomoldic, drusy. Quartz -subhedral, drusy. Clay-skeletal dolomoldic.

33

Feet below top RESIDUE: Chert-fine granular, white, abundant to skeletal dolomoldic, quartzose, drusy; extremely fine granular, white to buff stained. Quartz- anhedral, clear and subhedral, coarsely drusy. Clay-small amount skeletal dolomoldic. Dolomite-medium fine to fine, light gray, 90 percent. Chert-smooth and vory fine granular, dolomitic, 10 percent..... Residue: Chert-smooth, white, scattered dolomoldic, very fine granular white, quartzose and drusy, scattered dolomoldic. Quartz-subhedral, drusy. Dolomite-medium, very light gray..... 295 - 300RESIDUE: Quartz- small amount anhedral, finely granular, some subhedral, drusy. Clay -small amount fragmental dolomoldic. to white. Dolomite-medium coarse, light gray. Chert-small amount granular, white, slightly translucent, in part colluic, in part abundant dolomoldic. Quartz-subhedral, drusy, dolomoldic. Dolomite-medium fine, light gray and some pink coloration. Chert-small amount granular, grayish white, oolitic...... RESIDUE: Chert-fine granular, grayish white, fine to medium onlitic, few dolomoldic fragments. Quartz-subhedral, drusy, in part dolomoldic. Clay-fragmental skeletal dolomoldic. Dolomite-medium, light gray, sparingly cherty..... -----RESIDUE: Chert-smooth, white, scattered dolomoldic, very fine granular, abundant to skeletal dolomoldic, in part finely drusy; chalky, white to grayish. Quartz-small amount subhedral, drusy. Clay-finely fragmental dolomoldic. Dolomite-fine to medium fine, slightly buff-gray, 90 percent. Chert-grayish white, varie-320 - 325spherical concrctionary structure, small amount finely drusy, small amount very finely oolitic. Quartz--subhedral grains. Dolomite---fine, slightly buff-gray. Chert---small amount granular, grayish white, oolitic....325-330 RESIDUE: Chert-fine granular, grayish white, finely oolitic, few scattered dolomolds, slightly drusy. Clay--sponge-like, white and buff, finely fragmental. Dolomite-medium fine, gray. Chert-traces fine granular 330–335 RESIDUE: Chert-granular, white, abundant dolomoldic, quartzose and drusy. Quartz subhedral grains, drusy clusters. Silt-quartz grains, probably chemical in origin. Dolomite-- medium fine, light gray and fine, dark gray. Chert--some grayish white, fine335–340 cretionary structure. Quartz-subhedral, drusy, dolomoldic. Dolomite-medium fine, gray. Chert-some gray granular..... 340-345 RESIDUE: Chert-finely granular, abundant to skeletal dolomoldic; granular, light brown, oolitic, smooth, white. Quartz-coarse, subhedral grains and drusy clusters. Dolomite-fine to medium fine. Chert-small amount gray, granular 345--350 RESIDUE: Chert--fine granular, buff stained, skeletal dolomoldic; fine granular, grayish, smooth, white. Quartz-subhedral grains and drusy. ... 350–355 quartzose and drusy in part; smooth, white. Quartz-subhedral grains and druse. Dolomite-fine to medium fine, gray 355-360 RESIDUE: Chert-granular, slightly translucent-grayish, scattered to abundant dolomoldie, slightly drusy. Quartz-subhedral grains, few drusy clusters. Dolomite-medium fine, gray to buff-gray. Chert-small amount fine granular, light drusy. Quartz-small amount subhedral grains and druse. Dolomite-medium fine, gray and dark gray365--370 RESIDUE: Chert-fine granular, grayish; smooth, white, chalky, hard and rongh; smooth, buff stained, skeletal dolomoldic. Quartz-anhedral, granular, coarsely drusy; subhedral grains and druse.

Feet below top
Dolomite—fine, buff-gray and gray
Dolomite-fine to very fine, brown. Chert-small amount granular, grayish, pseudo- politic
RESIDUE: Chert—fine granular, grayish, slightly translucent, in part collic and pseudo-collitic, and in part drusy, with scattered dolomolds; smooth, grayish white. Quartz —subhedral grains and drusy flakes.
Dolomite—fine to very fine, gray and buff-gray
Dolomite—very fine, gray RESIDUE: Chert—very finely granulated, grayish white. Clay—flaky, dolomoldic
Dolomite—very line to fine, gray
Dolomite—fine to medium fine
Dolomite—very fine, gray. Chert—small amount granular, grayish
Dolomite—fine to medium fine
Dolomitefine and medium fine, buff and gray410-415 RESIDUE: Chertfine granular, white and buff, drusy, skeletal dolomoldic. Quartz
Dolomite medium, gray and slightly buff-gray
Dolomite—very fine, light buff and medium fine, light gray
Dolomite—very fine, buff, some gray
Dolomite—very fine to fine, buff-gray
Dolomite—medium to medium fine, light gray
Dolomite—very fine to medium fine, buff and gray
Dolomite—fine, light gray. RESIDUR: Quartz—very coarse, subhedral grains, drusy clusters; very finely granulated.
Dolomitemedium, light gray
Dolomite—medium, light gray, slightly porous

TANYARD SECTION, BURNET COUNTY, TEXAS

	Fe beloi	eet w to	
Tanyard formation, Staendebach member, 0-336 feet-			
Limestone- extremely fine, dove-gray. Chert-trace granular, white, oolitie RESIDUE: Chert-smooth, white, with scattered ooliths; some very fine granular, gra chert, slightly quartzose. Quartz-few banded, drusy fragments, subhedral grains.	ayish)	5
Limestone—extremely fine, dove-gray RESIDUE: Clay—sponge-like and flaky, white and maroon.		i_ 1	-
Limestone—extremely fine, very light gray RESIDUE: Chert—chalky, white, brittle. Clay—flaky, white.	10)— J	ι5
Dolomite—coarse, light gray, with maroon coloration, 60 percent. Limestone—cxtre fine, very light gray, 40 percent RESIDUE: Chert—chalky, soft, porous, white, slight maroon coloration. Clay—f white and maroon.			
Dolomite—coarse, light gray RESIDUE: Clay—finely fragmental dolomoldic and flaky, white and marcon.	20) - 2	25
Dolomite—coarse to very coarse, light gray			
Dolomite—coarse to very coarse, light gray RESIDUE: Clay—flaky and finely fragmental dolomoldic, white, gray, and maroon.	30	⊢ 3	35
Dolomite—very coarse, light gray, in part marcon stained. RESIDUE: Clay—small amount flaky and finely fragmental dolomoldic, marcon white.	35 and	- 4	10
Dolomite—medium coarse, light gray, with maroon staining. RESIDUR: Chert—granular, gray to grayish white, quartzose in part. Quartz-s amount finely drusy. Clay—sponge-like and flaky, white.	mall		
Dolomite—medium to coarse, light gray RESIDUE: Clayvery small amount sponge-like, grayish.			
Dolomite—medium coarse to coarse, light pinkish gray RESIDUE: Chert—very fine granular, white, abundant dolomoldic. Clay—small am sponge-like, dolomoldic, and flaky.	50 ount	- 5	5
Dolomite—coarse, light gray, in part slightly pinkish RESIDUE: Chert—very fine granular to slightly chalky, white, skeletal dolomo Quartz—subhedral grains. Clay—finely fragmental, skeletal dolomoldic.	55 Idic.	- 6	50
Dolomite—coarse, in part pinkish, light gray RESIDUE: Quartz—fine to coarse, subhedral grains, and dolomoldic, drusy clus Claywhite and maroon, finely fragmental dolomoldic.	60 sters.	- 6	55
Dolomite—medium coarse to coarse, some fine, pinkish and light gray RESIDUE: Clay—small amount, white, some marcon, finely fragmental dolomo Chert-trace soft chalky.	65 Idic.	- 7	0
Dolomite—coarse, light gray, and slightly pinkish. RESIDUE: Chert—very fine granular, some smooth, white, finely drusy, sparingly of moldic. Quartz—very fine to medium, subhedral grains, drusy clusters. Clay—sponge and finely fragmental dolomoldic, gray to white.	lolo- -like		
Dolomitc-fine to medium fine, light gray, some pink RESIDUE: Chert-very fine granular, scattered to abundant dolomoldic, white, s maroon. Quartz-highly drusy, dolomoldic flakes and fragments; fine to very co. subhedral grains.	ame	- 8	0
Dolomite—medium, light gray with pinkish streaks RESIDUE: Quariz—anhedral, slightly drusy; fine, angular and subhedral grains. Cl. very small amount fragmental dolomoldic.	ау—		
Dolomite—medium fine to medium, gray RESIDUE: Chert—chalky, soft, white, skeletal dolomoldic; trace smooth, white. Ch white, finely fragmental dolomoldic. Spicules—few extremely fine, needle-like.			
Dolomite—medium to modium fine, light gray, slightly pinkish gray and gray RESIDUE: Clay—very small amount sponge-like and finely fragmental dolomo · Spicule—long, needle-like.	ldic.		
Dolomitemedium fine to medium, gray. RESIDUE: Chertvery fine granular, gray, scattered to abundant dolomoldic; cha white, abundant to skeletal dolomoldic; chalcedonic, in part drusy. Claywhite, fi fragmental, dolomoldic. Quartzdrusy fragment, subhedral grains.	95 ulky, nely	-10	0

Feet below top ...100--105 Dolomite-coarse, light gray RESIDUE: Chert-chalky, white, abundant to skeletal dolomoldic, in part finely fragmental. _____105-110 Dolomite-coarse, some medium fine, pinkish gray.... RESIDUE: Clay-small amount sponge-like, flaky, white. Dolomite-coarse, light gray, some pinkish gray RESIDUE: Chert-chalky, grayish white to white, skeletal dolomoldic; smooth, white, -110-115skeletal dolomoldic; fine granular, white, drusy, in part dolomoldic. Quartz-anhedral, granular, finely drusy, subhedral grains. Clay-white, finely fragmental dolomoldic. Dolomite-medium, light gray..... RESIDUE: Chert-very fine granular, white, skeletal dolomoldic, in part finely drusy; chalky, grayish, skeletal dolomoldic. Quartz-very fine, subhedral grains. Clay-white, finely fragmental dolomoldic. Dolomite—medium to coarse, light and pinkish gray RESDUE: Chert—smooth, white to pinkish, very finely drusy, skeletal dolomoldic. Quartz—very finely drusy fragments. Clay—white, finely fragmental dolomoldic.120 - 125125 - 130Dolomite-coarse, light gray..... RESIDUE: Chert-smooth, white, abundant to skeletal dolomoldic; finely granular, white to buff, in part quartzose and finely drusy, skeletal dolomoldic. Quartz-few finely drusy flakes. Clay-white, finely fragmental dolomoldic.130-135 Dolomite-medium coarse, light gray..... RESIDUE: Chert-fine granular, white, quartzose and finely drusy, skeletal dolomoldic; smooth, white, skeletal dolomoldic. Quartz--anhedral, granular, drusy, dolomoldic, very fine, angular to medium, subhedral grains. Clay-white, finely fragmental, dolomoldic. Silt---clear quartz, probably chemical in origin. Dolomite-medium coarse, light gray RESIDUE: Quartz--fine to very coarse, subhedral grains, few drusy clusters, Chert-very fine granular, white, abundant to skeletal dolomoldic, quartzose and drusy. Clayvery finely fragmental, dolomoldic, white and pink.140-145 Dolomite-medium, light gray..... RESIDUE: Quartz-anhedral, dolomoldic, drusy, subhedral grains and clusters. Chert -small amount chalky, dolomoldic. Clay-cream and white, very finely fragmental dolomoldic. Dolomite-medium, light gray and fine, buff-gray, 90 percent. Chert-very fine granular,145-150 white, 10 percent. RESIDUE: Chert-white, variegated smooth and fine granular, oolitic; very fine granular, cream-white, abundant to skeletal dolomoldic, drusy in part. Quartz-fine to coarse, subhedral grains, drusy clusters. Clay-finely fragmental dolomoldic, white. 150-155 Dolomite-fine to medium gray, in part slightly pinkish. RESIDUE: Chert-very fine granular, white and pink, drusy, abundant to skeletal dolo-moldic. Quartz-fine to very coarse, subhedral grains, drusy clusters. Clay--white, finely fragmental dolomoldic. Dolomite-fine to medium, gray...... RESIDUE: Quartz-fine to coarsely drusy fragments; fine to very coarse, subhedral155-160 grains. Chert--fine granular, grayish white, scattered to abundant dolomoldic. Dolomite-medium fine to medium, gray..... RESIDUE: Quartz-drusy flakes and fragments, in part dolomoldic; fine to coarse, subhedral grains. Clay-white and maroon flecked, skeletal dolomoldic. Dolomite-medium, gray and fine, buff-gray.....165–170 RESIDUE: Chert-fine granular, white, some pink, abundant to scattered dolomoldic. Quartz-small amount finely drusy. Clay-white, finely fragmental dolomoldic. Dolomite--medium to fine, gray..... ...170–175 RESIDUE: Chert-fine granular, white, quartzose and drusy, abundant dolomoldic. Quartz—very fine to coarse, subhedral grains. Dolomite-medium fine, buff-gray....175–180 RESIDUE: Chert-very fine granular, grayish white, partly quartzose and drusy, abundant dolomoldic. Quartz-very fine to medium, subhedral grains. Dolomite-medium to medium coarse, light gray..... RESIDUE: Clay-small amount, grayish, finely fragmental dolomoldic. ...185--190 Dolomite-medium, light gray..... RESIDUE: Clay-small amount sponge-like, gray and white, finely fragmental dolomoldic.

Feet below top Dolomite-medium to fine, light gray, 90 percent. Chert-fine granular, white, 10 percent RESIDUE: Chert-granular, white, scattered dolomoldic. Quartz-very fine to coarse, subhedral grains, drusy, dolomoldic fragments. Dolomite-fine, slightly pinkish gray, 90 percent. Chert-grayish white, granular, colitic, 195–200 - - . 10 percent pseudo-oolitic, in part quartzose; smooth, white, oolitic. Quartz-very fine to coarse, subhedral grains, Dolomite-medium fine, pink, 90 percent. Chert-white, porcelaneous, 10 percent ...200-205 RESIDUE: Chert-ordinary smooth and porcelaneous, white; very fine granular, white to pink, scattered dolomoldic. Quartz-fine to coarse, subhedral grains. Clay-small amount white and pink, very finely fragmental dolomoldic. 205 - 210amount white and pink, finely fragmental and skeletal dolomoldic. Dolomite-medium fine, gray, some pink staining. Chert-trace white, granular, very finely oolitic 2 RESIDUE: Chert-fine granular, grayish white and pink, in part abundant dolomoldic and quartzose, in part very finely oolitic. Quartz-finely drusy flakes, fine angular grains. Clay-small amount white, finely fragmental, Dolomite-fine, gray and pink, 80 percent. Chert-fine granular and porcelaneous, white, quartzose and drusy in part, 20 percent ..215-220 RESIDE: Chert-granular, white to gray, very scattered dolomoldic, in part colitic, with trace comolds, in part quartzose and drusy; ordinary and porcelaneous smooth, white to grayish chert. Quartz-anhedral, clear, drusy fragments; fine to very coarse, subhedral grains. Clay-small amount finely fragmental. Dolomite-fine to medium, slightly pinkish gray..... 220-225 RESIDUE: Chert-small amount grayish, finely graunlated. Quartz-few subhedral grains. Dolomite-medium to medium fine, gray and pinkish..... 225 - 230RESIDUE: Quartz-fine to coarse, subhedral grains; very finely granulated. Clay-small amount white, finely fragmental dolomoldic. Dolomite---medium coarse to medium, light gray..... RESIDUE: Clay-small amount white, finely fragmental dolomoldic. Dolomite-medium fine to fine, gray, some pinkish RESIDUE: Chert-very fine granular, white, abundant to skeletal dolomoldic. Quartz---......235--240 medium, subhedral grains, very fine, angular granules, few finely drusy flakes. Claywhite, finely fragmental, dolomoldic. Dolomite-fine to medium fine, pinkish gray...240-245 RESIDUE: Chert-grayish, finely granulated, rhombic granules; trace fine granular, white, dolomoldic. Clay-small amount white, finely fragmental dolomoldic. Dolomite-fine to medium coarse, light gray.....245 - 250RESIDUE: Quartz-finely granulated; trace dolomoldic, finely drusy; few subhedral grains. Clay-white, skeletal dolomoldic. Dolomite-fine, pinkish gray, some medium, light gray..... -250-255RESIDUE: Chert-very fine granular, very finely and abundantly drusy, skeletal dolomoldic. Quartz-very fine granules and fine subhedral grains. Dolomite-fine, slightly pinkish, some medium fine, light gray .255 - 260RESIDUR: Chert-fine granular, grayish white, in part very finely drusy, skeletal dolo-moldic; smooth, white, skeletal dolomoldic. Quartz-fine to medium, subhedral grains, and finely granulated. Dolomite-medium fine to coarse, light gray, porous.... RESIDUE: Quartz-very small amount very fine, subhedral grains. Dolomite-fine, pinkish gray______ RESIDUE: Chert-small amount grayish, finely granular. Quartz-few subhedral grains. Clay-small amount very finely fragmental, dolomoldic, Dolomite-fine to medium fine, some pinkish..... RESIDUE: Chert-granular, white, in part very finely dolomoldic; chalky, in part very finely porous and dolomoldic; very fine granules in part thombic. Clay-small amount sponge-like, white to buff.

Feet below top Dolomite-medium to fine, light gray and pinkish..... RESIDUE: Quartz-subhedral, very fine to very coarse grains. Chert-smooth, tan, and pink, skeletal dolomoldic, in part quartzose and drusy. Dolomite-medium, light gray and fine, pinkish gray..... RESIDUE: Chert—granular, white, finely drusy, skeletal dolomoldic. Quartz—very fine, angular grains. Dolomite-fine, pinkish, and medium coarse, light gray..... RESIDUE: Chert-granular, white, highly drusy, skeletal dolomoldic. Quartz-drusy, dolomoldic flakes and very finc, subhedral grains. Clay-gray, flaky, and very finely fragmental dolomoldic. Dolomite-medium coarse and fine, pinkish gray..... RESEDUR: Chert-small amount fine granular, white and pinkish, dolomoldic. Clay-small amount sponge-like, very fively fragmental and dolomoldic. 295–300 Dolomite-medium to coarse, light gray, some pinkish RESIDUE: Chert-small amount brittle, chalky. Clay-gray, fragmental dolomoldic. Dolomite-fine, gray, in part pinkish 300-305 Restours: Clay- small amount white, gray, pink, finely fragmental dolomoldic. Quartz -trace finely drusy. Spicules—few very finc, needle-like. ... 305-310 Dolomite-fine, buff-gray and pinkish RESIDUE: Chert-very fine granular, white to gravish, some pink, scattered to abundant, finely dolomoldic; smooth, white, with finely dolomoldic contact surfaces; chalky, slightly dolomoldic. Quartz-small amount dolomoldic, drusy; few subhedral grains. Clay-white, very finely fragmental dolomoldic. Dolomite-medium coarse to medium fine, light gray and pink 310--315 RESIDUE: Chert-fine granular, white and pink, in part drusy, abundant to skeletal dolomoldic; chalky, small amount dolomoldic. Quartz-dolomoldic, drusy fragments, fine, subhedral grains. 315–320 Dolomite-coarse, gravish white ... RESIDUE: Chert-granular, white, abundant to skeletal, coarsely dolomoldie, in part drusy. Quartz-fine to coarse subhedral grains, dolomoldic, drusy fragments. Dolomite-medium to medium fine, light gray, some pink 320 - 325RESIDUE: Chert--granular, white, quartzose and finely drusy, skeletal dolomoldic. Quartz-finely drusy, dolomoldic, few subhedral grains. Dolomite-medium to coarse, light gray RESIDUE: Chert-small amount finely granulated. Shale-green to dark green, waxy. Clay-small amount white, finely fragmental. Dolomite—coarse, light gray..... Respute: Shale—small amount light green, some dark green, waxy, in fine fragments. Chert-trace finely granulated. Tanyard formation, Threadgill member, 336-538 feet----Dolomite-coarse to very coarse, light gray, some pinkish..... RESIDUE: Shale-dark maroon and dark green, waxy, in part dolomoldic. Clay-small amount sponge-like and finely fragmental delomoldic, white. Dolomite-coarse, light gray, some buff staining..... RESIDUE: Shale-green, waxy, in very fine fragments. Chert-small amount soft chafky; trace granulated. Dolomite--coarse, light gray, some maroon stained 345 - 350RESIDUE: Shale-light green, some dark green and maroon, waxy. Chert-small amount granulated; trace smooth, lacy. Dolomite-coarse, light gray, some purplish..... 350-355 RESIDUE: Shale-waxy, grayish white to purplish and maroon, fragmental skeletal dolomoldic. Clay--small amount flaky. Quariz-few medium to coarse, subhedral grains. RESIDUE: Chert-small amount finely granulated. Shale-waxy, greenish and purplish to maroon. Clay-trace flaky. Dolomite--coarse to medium coarse, grayish white RESIDUE: Chert-granulated, white. Shale-small amount waxy, green and maroon, fragmental dolomoldic. Dolomite-coarse, very light gray, some slightly pinkish..... 365-370 RESIDUE: Shale-waxy, green, some maroon, traces of dolomolds. Chert-small amount white, granulated.

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Feet below top
Dolomite—coarse, very light gray
Dolomite—coarse, some fine, very light gray, some purplish staining, 60 percent. Lime- stoneextremely fine, grayish white, some purplish staining, 40 percent
Dolomite—coarse, white, 80 percent. Limestone—extremely fine, grayish white, 20 per- cent. Chert—trace gray granular
Dolomite—coarse to very coarse, grayish white
Limestone—extremely fine, grayish white. Dolomite—trace white
Dolomite—very coarse, white, trace pink
Dolomite—very coarse, light gray and pink, 50 percent. Limestone—extremely fine, grayish white, 50 percent
Limestone—extremely fine, slightly grayish white
Dolomite—very coarse, white. Limestone—small amount extremely fine, very light gray, dolomitic and dolomoldic
Limestone—extremely fine, grayish white, trace dolomitic, trace purplish color
Limestoneextremely fine, white, trace dolomitie 420-425 RESIDUE: Chert-granular, semi-transclucent-gray, highly porous to lacy; smooth, white, roughly concretionary nodules. Shalevery small amount waxy green.
Limestoneextremely fine, very light gray, traces dolomitic
Limestone-extremely fine, very light gray; trace finely dolomitic
Limestone—extremely fine, slightly grayish white, trace dolomitic
Limestone—extremely fine, slightly grayish white, trace dolomitie
Limestone—extremely fine, slightly grayish white445-450 RESIDUE: Shale—few fine, waxy, green flakes.
Limestone—extremely fine, white, in part with slightly greenish tint
Limestone—extremely fine, slightly gravish white
Limestone—extremely fine, grayish white
Limestone—extremely fine, grayish white
Limestone—extremely fine, light gray, trace coarsely dolomitic
Limestoneextremely fine, grayish white

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Feet below top Limestone-extremely fine, grayish white, in part with indistinct, finely fragmental (pseudo-oolitic?) structure RESIDUE: Quartz-anhedral, in fragments with rough, granulated surfaces. Chert-.....480-485 trace white, nodular, and granular, slightly translucent-grayish. .485-490 Limestone-cxtrcmely fine, light gray ... Response: Shale-very small amount waxy, drab green. Silt-small amount, rhombic grains. Limestone --extremely fine, grayish white, trace dolomitic...... RESIDUE: Chert-trace, fine granular. Silt-rhombic grains. RESIDUE: None. Limestone-extremely fine, slightly grayish white, in part coarsely dolomitic RESIDUE: Shale-very small amount waxy, green. Silt-trace rhombic grains. Limestone---extremely fine, grayish white, in part slightly dolomitic, with slightly greenish tint RESIDUK: Shale-small amount waxy, green, and brownish green. Chert-trace fine, silicified rhombs. Silt-rhombic grains. Limestone-extremely fine, grayish white, 50 percent. Dolomite-coarse, light gray, 50 percent . •_____ RESIDUE: Shale-trace dark green, waxy. Silt-trace rhombic grains. Dolomite-coarse, light gray..... RESIDUE: Quartz-anhedral, cloudy, rough to lacy fragments; trace finely granulated. Shale--small amount brownish green, waxy. Dolomite-medium to coarse, gray..... RESIDUE: Shale-drab green, some maroon, waxy. Silt-very fine, rhombic grains. Dolomite-- coarse to medium, light gray..... 525-530 RESIDUE: Shale-dark green, waxy, trace of dolomolds. Quartz-small amount finely granulated. Dolomite--coarse, light gray..... RESIDUE: Chert-small amount white, granulated, trace white, nodular. Clay-sponge-like, in part dolomoldic. Shale-small amount green, waxy. Silt-very fine, in part rhombic grains, Quartz-trace anhedral, lacy. Dolomite-medium, light gray, 60 percent. Limestonc-fine, purplish marcon, 40 per-cent RESIDUE: Ouartz-finely granulated, in part very slightly orchid tinted; trace anhedral, lacy. Chert-trace rough, white, chalky. Wilberns formation, 538-570 feet-Dolomite-very fine, dull purplish maroon, some buff-gray RESIDUE: Chert-very finely granulated, slightly purplish tinted. Clay-small amount very finely dolomoldic, white to buff. Dolomite-very fine, gray, some purplish coloration, in part siliceous..... RESIDUE: Chert-very fine granular, grayish white to carthy gray, very scattered to abundant finely dolomoldic; some smooth, white, slightly chalcedonic; trace chalky. Clay-gravish white, mostly extremely fine fragmental. Quartz-small amount drusy. Dolomite -very fine, buff-gray and slightly purplish, in part siliceous. Chert-some dolomoldic; some white, porcelaneous, some slightly chalcedonic. Clay-grayish white, extremely fine fragmental. Dolomite-very fine to extremely fine, buff-gray, some purplish, in part siliceous. Chert 555--560 dolomoldic; some buff stained and purplish, fine granular. Clay-grayish white, extremely fine fragmental. Dolomite--very fine, buff-gray, trace purplish. Chert- some very fine granular, white 560-565 RESIDUR: Chert-very fine granular, buff, pink stained, white, and slightly translucent-grayisb, non-dolomoldic to abundant dolomoldic; some chalky, skeletal dolomoldic. Quartz-some finely drusy flakes. Clay-gravish white, extremely fine fragmental. Dolomite-very fine, buff-gray and purplish gray565-570 RESIDUE: Clay-white, some pink, finely fragmental, skeletal dolomoldic. Chert-soft, chalky, scattered dolomoldic.

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